

How do millennial proxy reconstructions methods stack up?



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$$\begin{aligned}\frac{dX}{dt} &= -\sigma X + \sigma Y + f_0 \cos \theta \\ \frac{dY}{dt} &= -XZ + rX - y + f_0 \sin \theta \\ \frac{dZ}{dt} &= xy - bZ,\end{aligned}$$

$$\sigma = 10, r = 28, b = 8/3$$

Allen and Stott, 2003, Climate Dynamics
Section 4

Distribution of 1-day averages

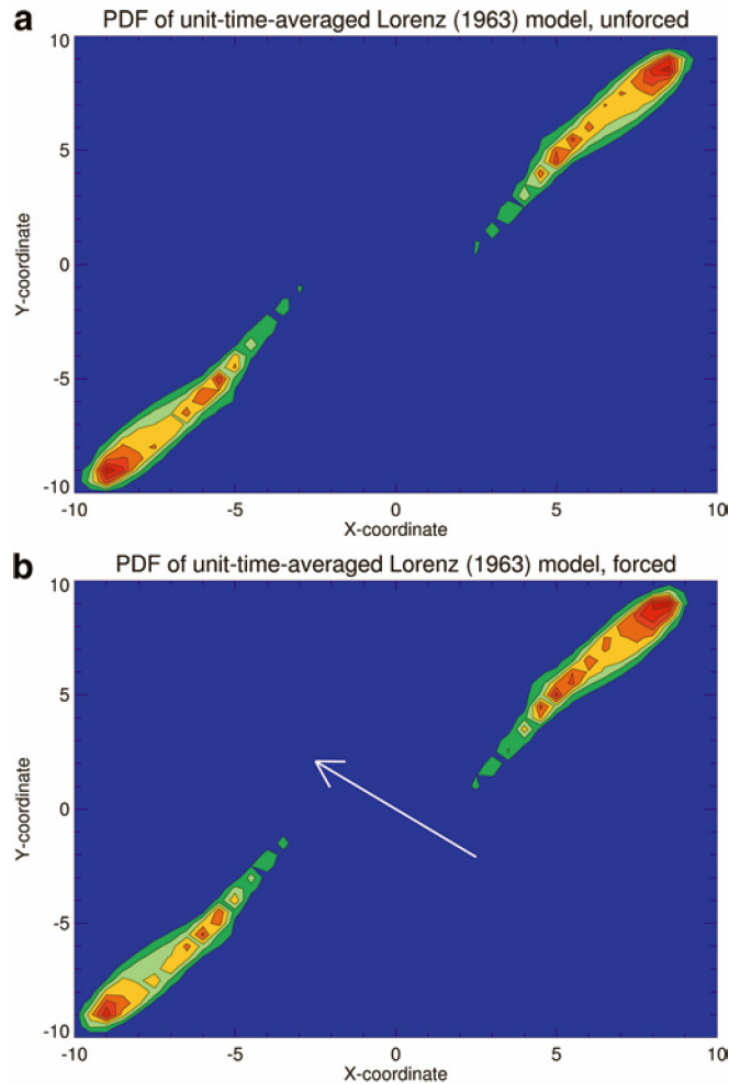


Fig. 3 a: estimate of the attractor probability density function (PDF) of the unforced Lorenz (1963), system. The plot shows a two-dimensional histogram of the location of “one-Lorenz-day” time-averaged values of the (X, Y) variables obtained from a long integration. **b:** PDF after imposing a steady forcing in the (X, Y) plane in the direction shown by the *arrow*, following Palmer (1999)

Distribution of 500-day averages

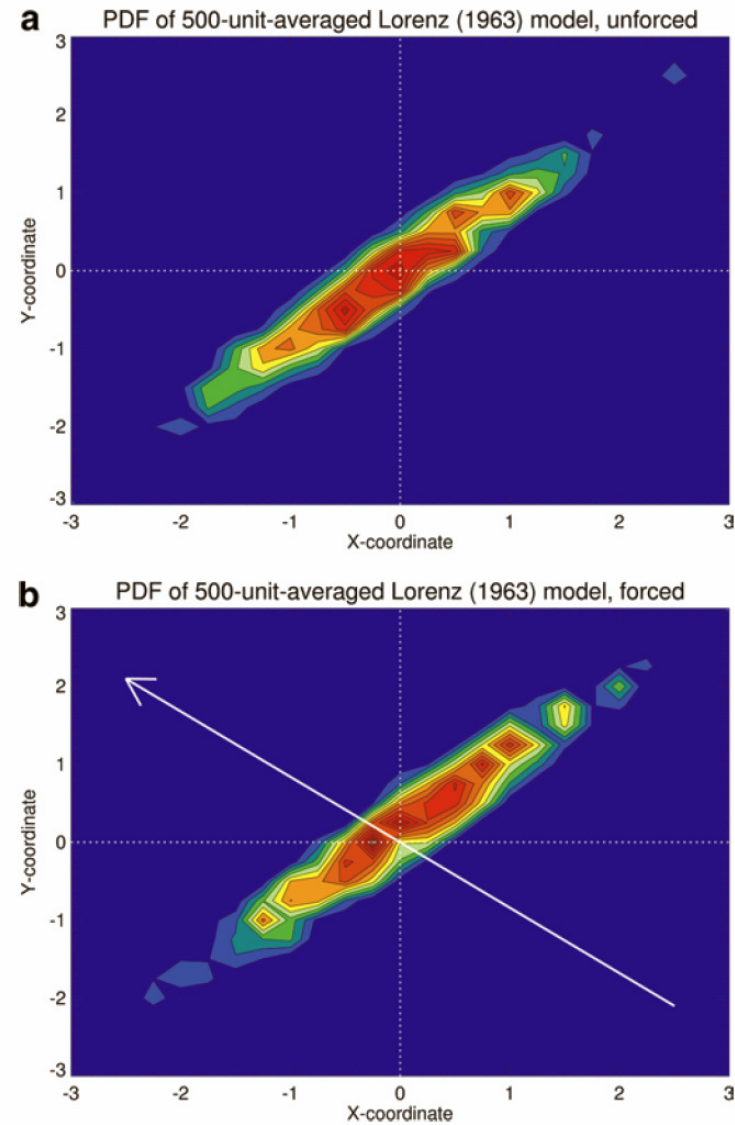
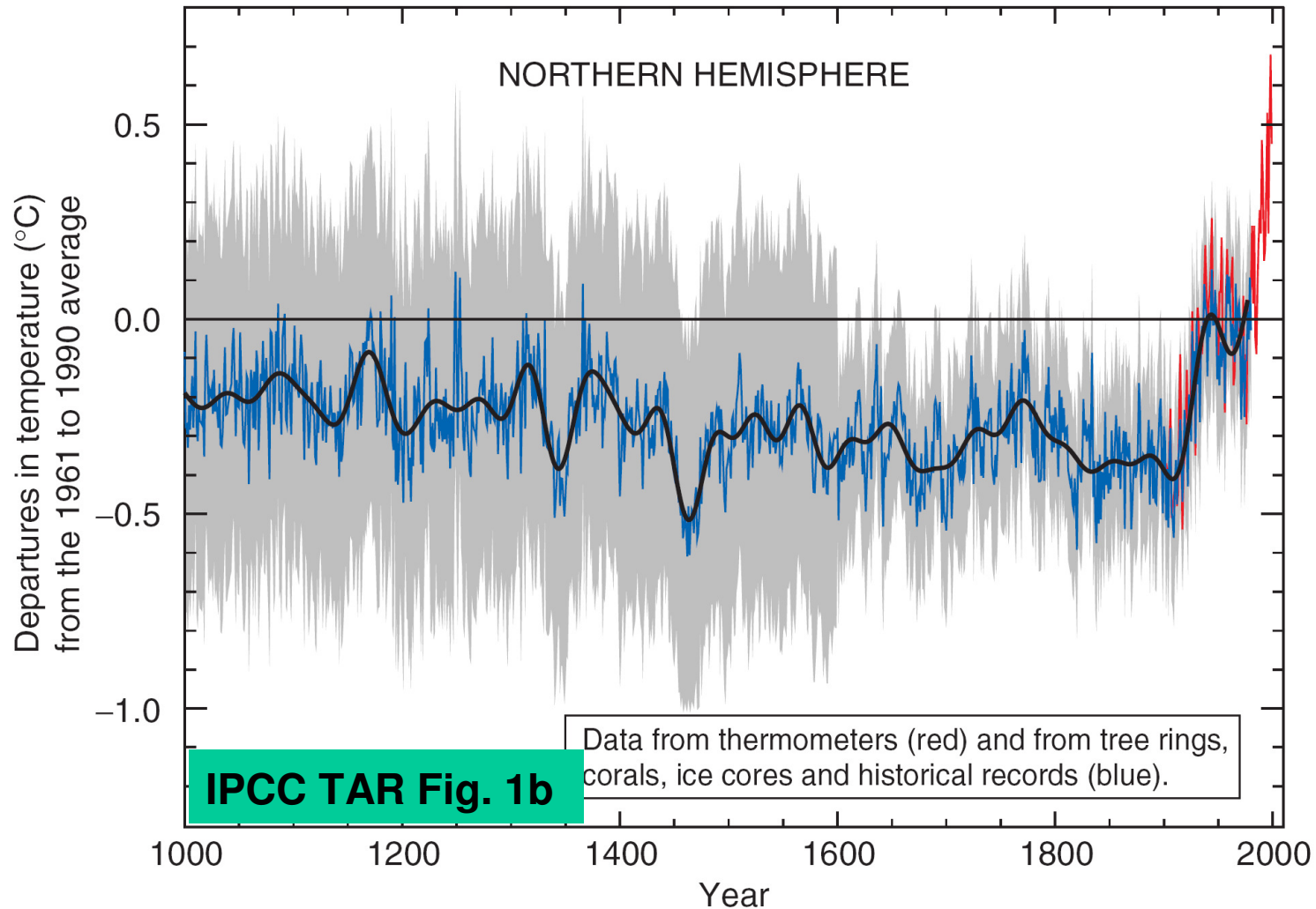


Fig. 4a, b As Fig. 3, but based on 500-Lorenz-day averaged data, to show the impact of time-averaging on the distributional properties of variability generated by a chaotic system: a simple consequence of the Central Limit Theorem

Outline

- **Background**
- **Reconstruction calibration techniques**
- **Their evaluation**
- **Conclusions**
- **Two extensions**
- **Comments??**

The “hockey stick” (TAR SPM)

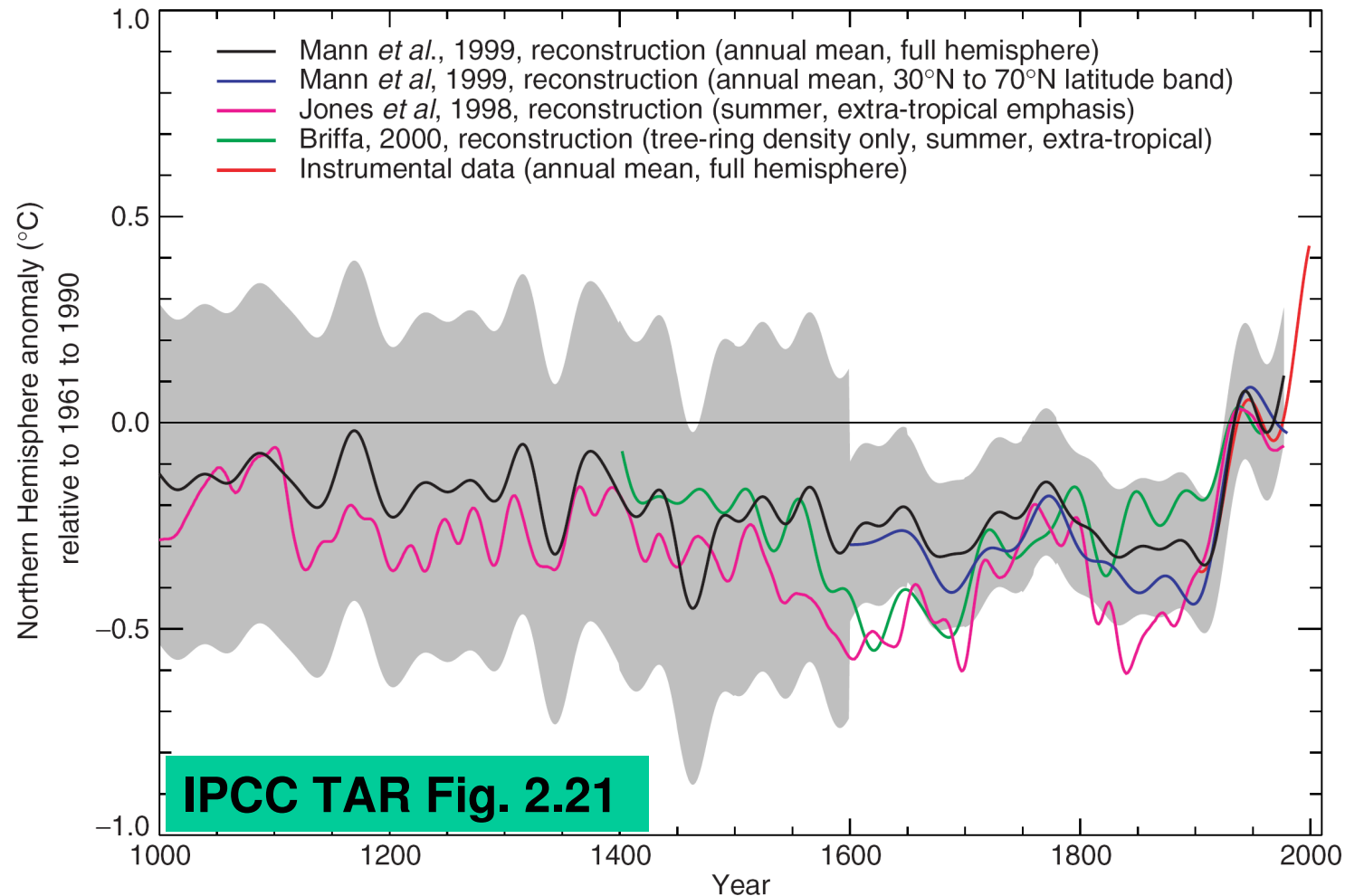


- “It is *likely* that the 1990s have been the warmest decade and 1998 the warmest year of the millennium”

The “hockey-stick” debate ...

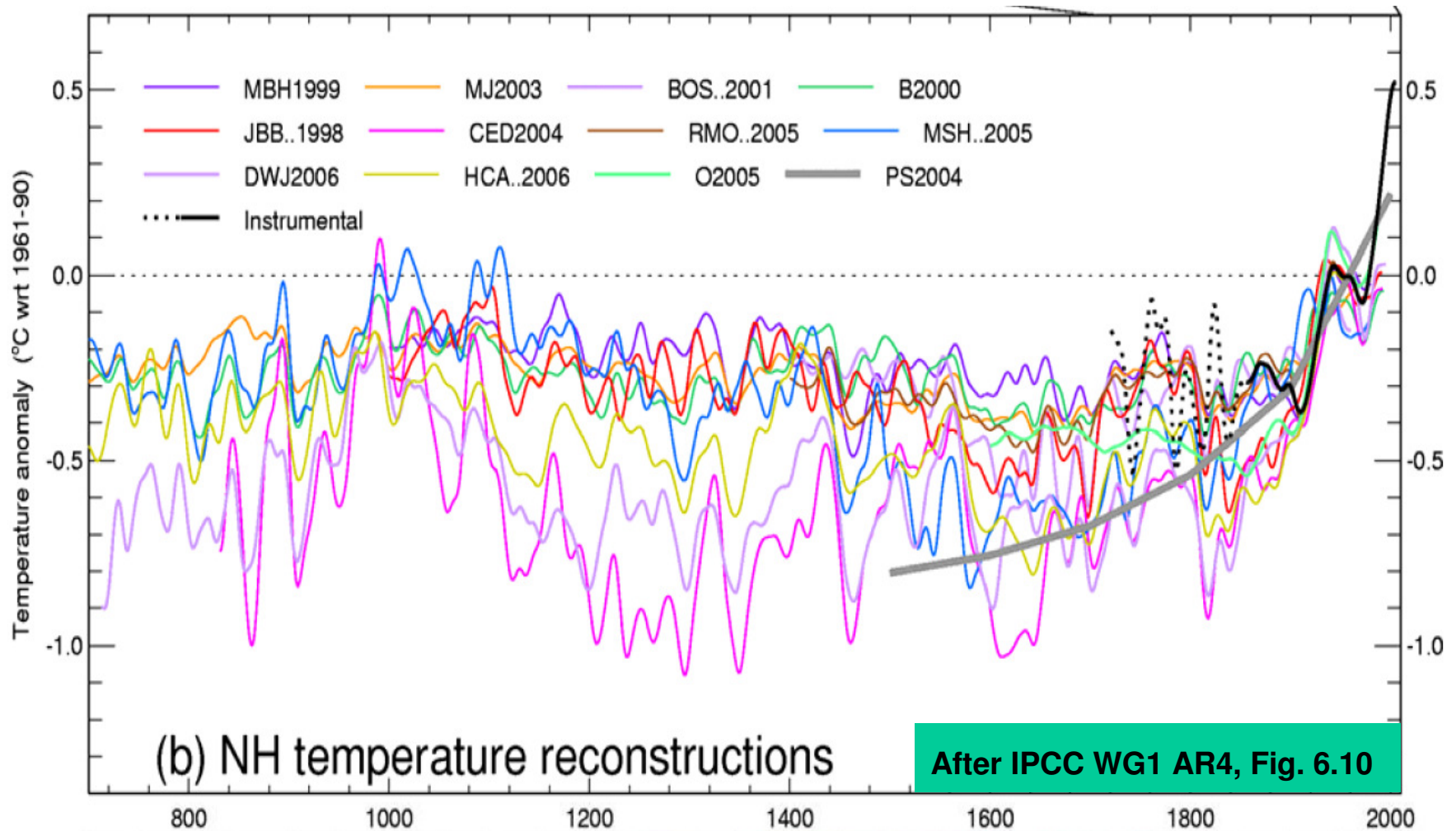
- Iconic figure and bold statement drew lots of interest from the sceptics
- Some Canadians particularly vocal
- Criticism focused on
 - Statistical method (influence of centering on EOFs)
 - Choice of proxies (use of Bristlecone Pines)

There was another figure in the TAR ...



- Indicating that the general form was robust,
- and that some reconstructions were more variable than others

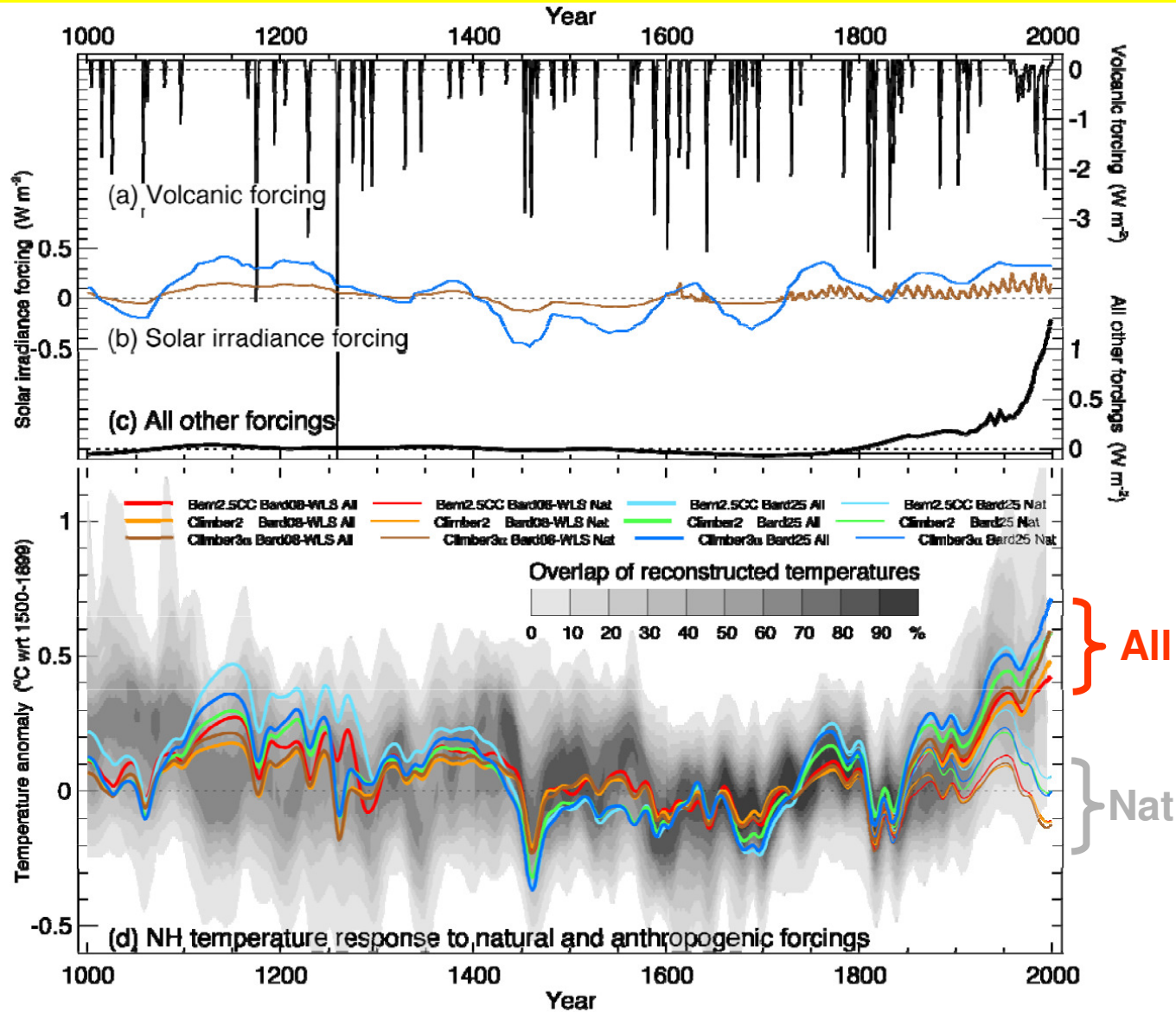
Reconstructions at time of AR4



Interpreting the reconstructions with models

- There now exist several long climate simulations with reconstructed volcanic, solar and GHG forcing
 - With AOGCMs (ECHO-G, CSM 1.4, HadCM3 ...)
 - EMICS
 - EBMS
 - (Note – now even more as a consequence of CMIP5)
- These seem able to reproduce some of the features seen in the reconstructions...
- Detection studies confirm the presence of natural and anthropogenic signals...

Temperature of last millenium



After IPCC WG1 AR4, Fig. 6.14

IPCC AR4 Assessment

- Average Northern Hemisphere temperatures during the second half of the 20th century were ***very likely*** higher than during any other 50-year period in the last 500 years and ***likely*** the highest in at least the past 1300 years.
- Assessment refers to longer time scales than in the TAR
- Many more reconstructions available
- Some detection results indicating external influence, detectable in all reconstructions that were examined
- Some inter-comparison of reconstruction techniques
- We have now undertaken a more complete assessment of available techniques (although still far from exhaustive)

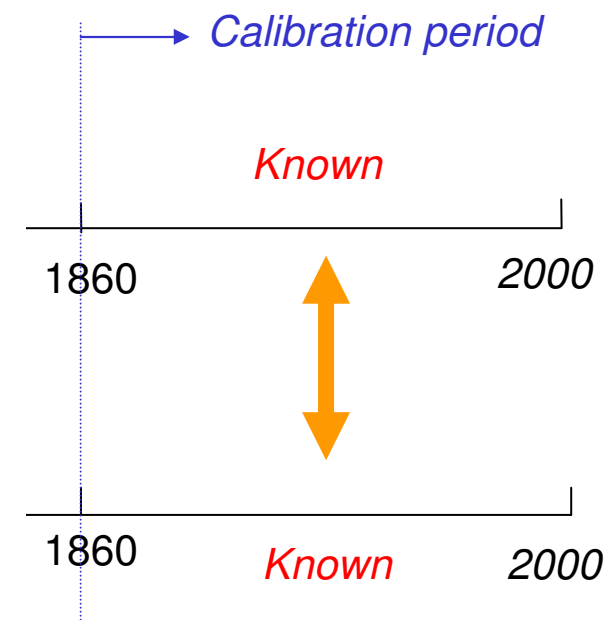
The reconstruction techniques



Photo: F. Zwiers

The Reconstruction Problem

Training period



Identify a statistical relationship between a collection proxies and NH temperature

Two types of reconstruction techniques

- **CPS – composite plus scale**
 - Average (or composite) proxies into some index (e.g., just average, and make dimensionless)
 - Calibrate the composite to hemispheric mean temperature from instrumental data
- **CFR – climate field reconstruction**
 - EOF regression, or other technique, to reconstruct hemispheric temperature field
 - Used, for example, to reconstruct SSTs back into 1800's using sparse instrumental data
 - Spatially average the reconstructed field to estimate hemispheric mean temperature

Reconstruction techniques

CPS	• Ordinary Least Squares	$T_t = \beta P_t + \varepsilon_t$
	• Total Least Squares	$T_t = \beta (P_t - \eta_t) + \varepsilon_t$
	• Variance Matching	$T_t = \beta P_t + \varepsilon_t$ $\hat{\beta} = [V(T) / V(P)]^{1/2}$
	• Inverse Regression	$P_t = \beta T_t + \varepsilon_t$
	• Kalman Filter/Smoothen	$P_t = \beta T_t + \varepsilon_t$ $T_t = \varphi T_{t-1} + \delta F_t + \omega_t$
CFR	• MBH (1998)	$T_t = \int_{NH} \sum_k \lambda_k u_{t,k} v_k dA + \varepsilon_t$
	• RegEM	

Standard parameter estimation approach (STD)

- Maximum likelihood of state process given proxies and estimated responses to forcing.

$$\ell_{\mathbf{P}}(\Theta) = -0.5 \sum_{t=1}^N \ln |\mathcal{A}\mathbf{S}_{t|t-1}\mathcal{A}^T + \mathcal{R}|$$

$$- 0.5 \sum_{t=1}^N (\mathbf{P}_t - \mathcal{A}\mathbf{T}_{t|t-1})^T (\mathcal{A}\mathbf{S}_{t|t-1}\mathcal{A}^T + \mathcal{R})^{-1} (\mathbf{P}_t - \mathcal{A}\mathbf{T}_{t|t-1})$$

Kalman filter estimates:
non-linear function of parameters

Proxy data (\mathbf{P}_t)



State process:

NH temperature (\mathbf{T}_t)



Estimated influence
of external forcings (\mathbf{F}_t)



1000

1850

2007

Improved parameter estimation technique (ALL)

$$\begin{aligned}
 -2 \ell_{\mathbf{P}, \mathbf{T}}(\Theta) = & \sum_{t=1}^n \left[\ln |\mathcal{A}\mathbf{S}_{t|t-1}\mathcal{A}^T + \mathcal{R}| + (\mathbf{P}_t - \mathcal{A}\mathbf{T}_{t|t-1})^T (\mathcal{A}\mathbf{S}_{t|t-1}\mathcal{A}^T + \mathcal{R})^{-1} (\mathbf{P}_t - \mathcal{A}\mathbf{T}_{t|t-1}) \right] \\
 & + m \ln |\mathcal{R}| + \sum_{t=n+1}^N (\mathbf{P}_t - \mathcal{A}\mathbf{T}_t) \mathcal{R}^{-1} (\mathbf{P}_t - \mathcal{A}\mathbf{T}_t)^T \\
 & + (m-1) \ln |\mathcal{Q}| + \sum_{t=n+2}^N (\mathbf{T}_t - \phi \mathbf{T}_{t-1} - \delta \mathbf{F}_t) \mathcal{Q}^{-1} (\mathbf{T}_t - \phi \mathbf{T}_{t-1} - \delta \mathbf{F}_t)^T \\
 & + \ln |\phi \mathbf{S}_{n|n} \phi^T + \mathcal{Q}| \\
 & + (\mathbf{T}_{n+1} - \phi \mathbf{T}_{n|n} - \delta \mathbf{F}_{n+1}) (\phi \mathbf{S}_{n|n} \phi^T + \mathcal{Q})^{-1} (\mathbf{T}_{n+1} - \phi \mathbf{T}_{n|n} - \delta \mathbf{F}_{n+1})^T.
 \end{aligned}$$

Kalman filter estimates

Proxy data (\mathbf{P}_t)



State process:

NH temperature (\mathbf{T}_t)



Estimated influence
of external forcings (\mathbf{F}_t)



1000

1850

2007

simple

Competing parameter estimation technique (CAL)

- Uses only calibration period data.

$$-2 \ell_c(\Theta) = \sum_{t=n+1}^N [\ln \mathcal{R} + (\mathbf{P}_t - \mathcal{A}\mathbf{T}_t)^2 / \mathcal{R}] + \sum_{t=n+2}^N [\ln \mathcal{Q} + (\mathbf{T}_t - \phi\mathbf{T}_{t-1} - \delta\mathbf{F}_t)^2 / \mathcal{Q}]$$

- MLEs can be solved explicitly, for example,

$$\hat{\mathcal{A}} = \left(\sum_{t=n+1}^N \mathbf{T}_t^2 \right)^{-1} \sum_{t=n+1}^N \mathbf{T}_t \mathbf{P}_t \quad \hat{\mathcal{R}} = m^{-1} \sum_{t=n+1}^N (\mathbf{P}_t - \hat{\mathcal{A}}\mathbf{T}_t)^2$$

Proxy data (\mathbf{P}_t)



State process:

NH temperature (\mathbf{T}_t)



Estimated influence
of external forcings (\mathbf{F}_t)



1000

1850

2007

Comparing estimator properties

- Likelihood too complicated to derive asymptotic properties in all cases → used Monte-Carlo
- Generate data under assumed model:

$$P_t = A T_t + e_t, \quad e_t \sim N(0, R)$$

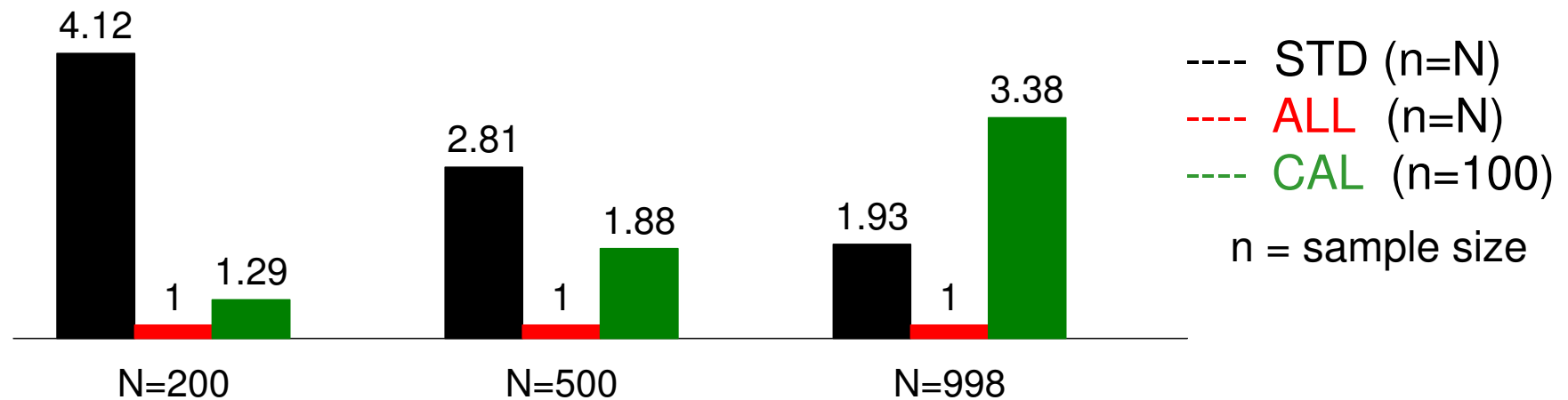
$$T_t = \phi T_{t-1} + \delta F_t + w_t, \quad w_t \sim N(0, Q)$$

- Specified reasonable values for A , ϕ , δ , R and Q .
- Used F_t from a simple EBM of the climate.
- Generated noise terms e_t and w_t from normal.

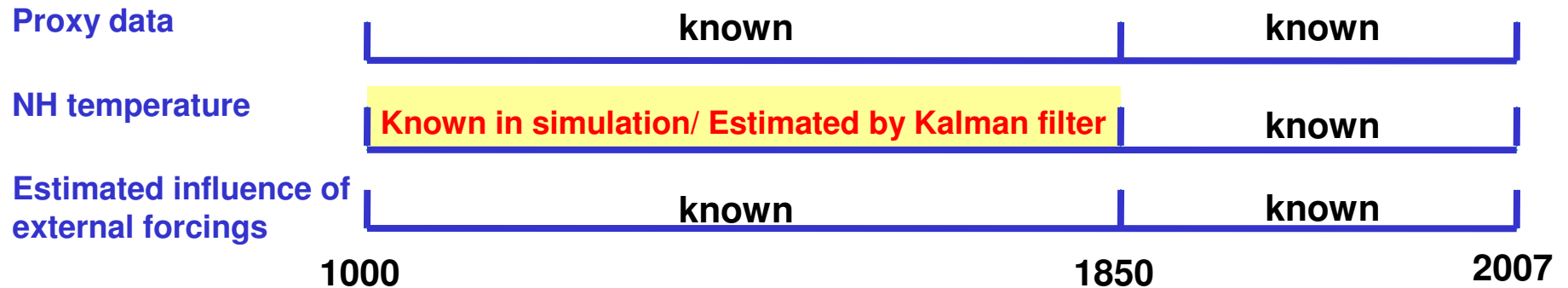
Relative efficiency of persistence parameter

Smaller value in efficiency = smaller estimator variance
= better estimator

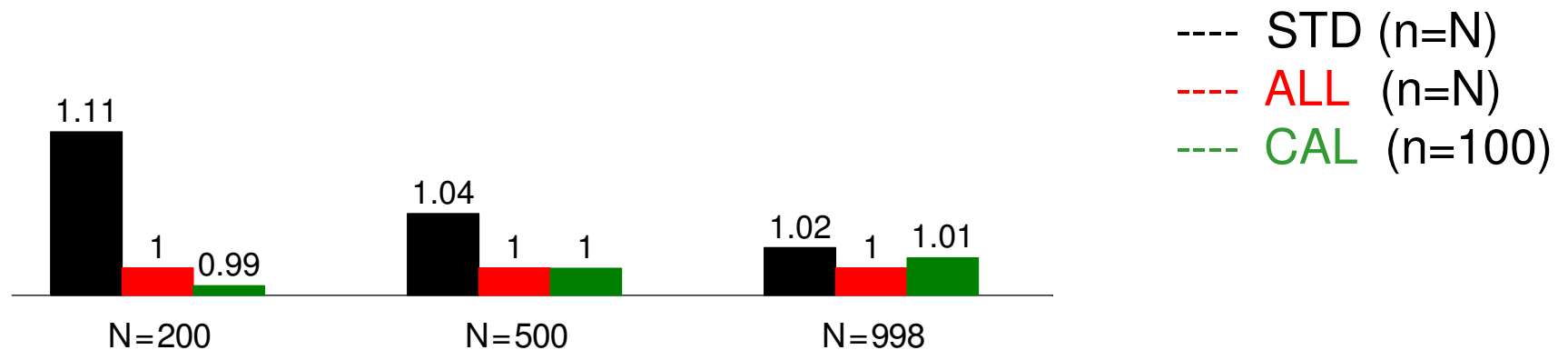
Estimated efficiency for ϕ (relative to **ALL**)



Reconstruction error



Reconstruction error (relative to ALL)



Smaller error = better approach

Robustness

$$e_t = \alpha e_{t-1} + \kappa_t, \quad \kappa_t \sim N(0, R)$$

$$P_t = A T_t + e_t,$$

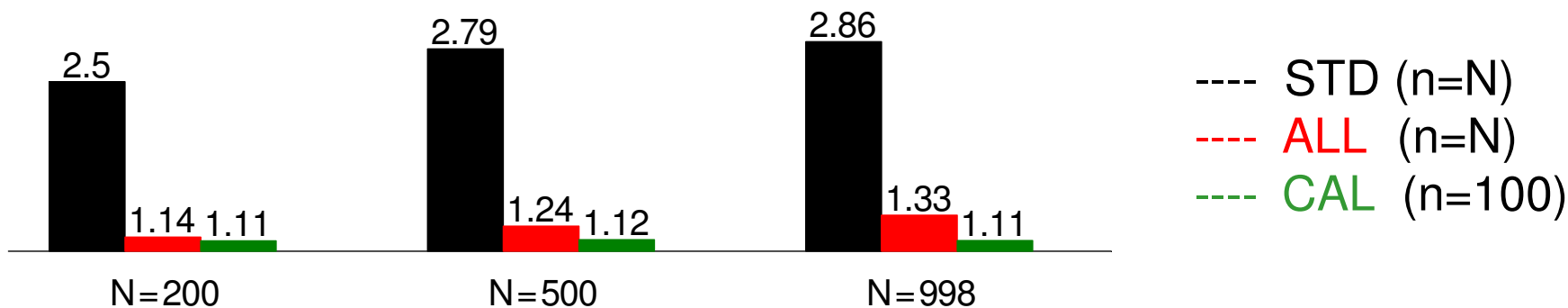
~~$$e_t \sim N(0, R)$$~~

$$T_t = \phi T_{t-1} + \delta F_t + w_t,$$

Estimators from CAL are asymptotically unbiased!

Change in reconstruction error
(relative to iid case)

Smaller change = more robust



Summary

- STD approach should be avoided.
 - Larger bias.
 - Lower efficiency.
 - Less robust.
- CAL approach is more robust
 - Reconstruction errors similar to ALL approach
 - Resistant to misspecification of state equation
 - MLEs easily obtained
- ALL approach if desire is to estimate state equation parameters

Evaluating the reconstruction techniques



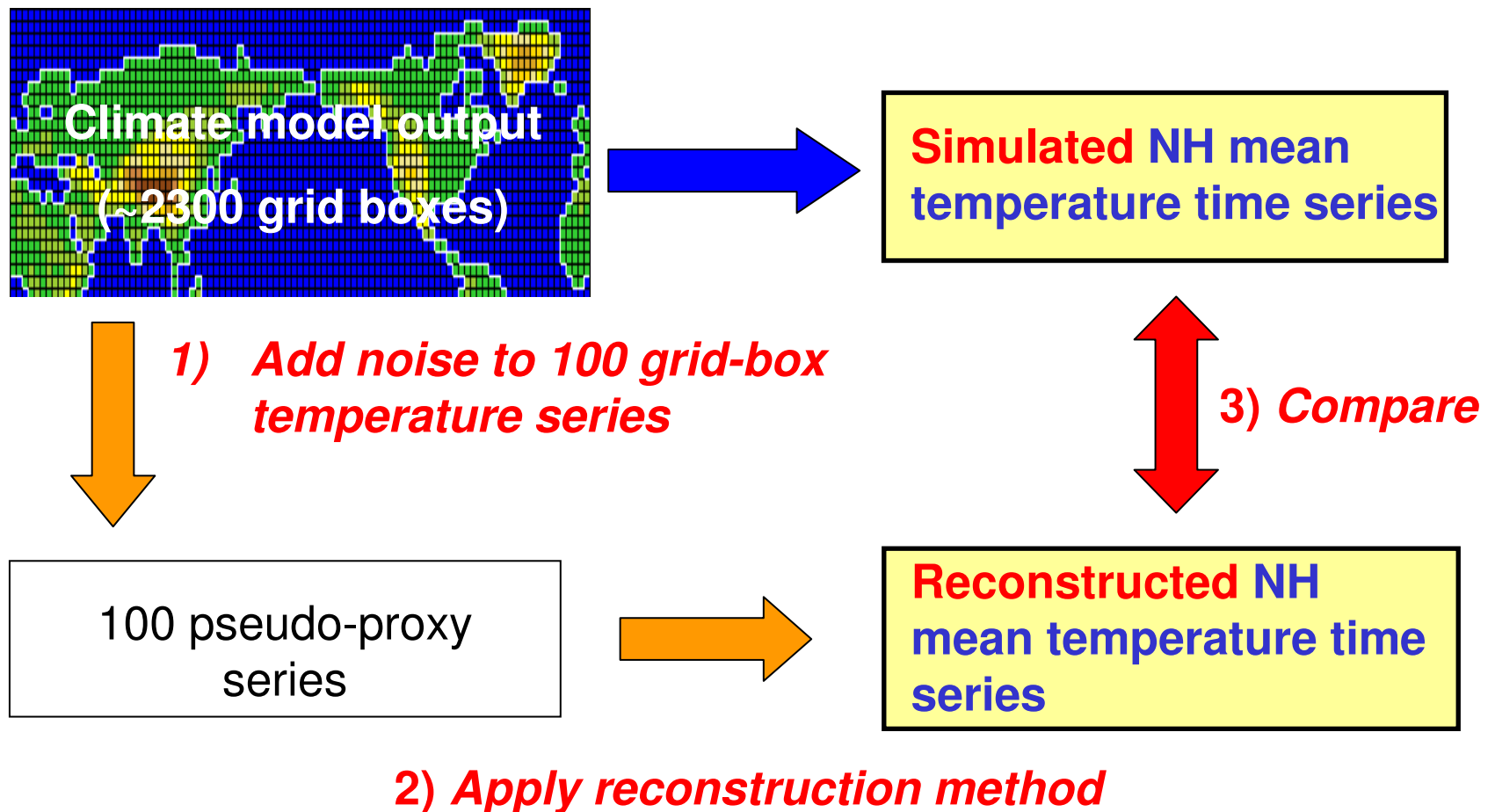
Photo: F. Zwiers

Evaluating Reconstruction techniques

- von Storch et al (2004) proposed using long climate simulations
- Idea was to
 - Sample model output at locations coinciding with proxies
 - Degrade by adding white noise
 - Apply reconstruction method
 - Compare with the known model simulated hemispheric mean temperature
- There are nuances to consider such as,
 - how much noise should be added?
 - what colour of noise should be added?
 - should one detrend prior to diagnosing the obs/proxy relationship?
 - what part of the spectrum should be used?
- The latter three questions are all broadly equivalent

Proposed evaluation technique

- Proposed by von Storch et al. (2004)

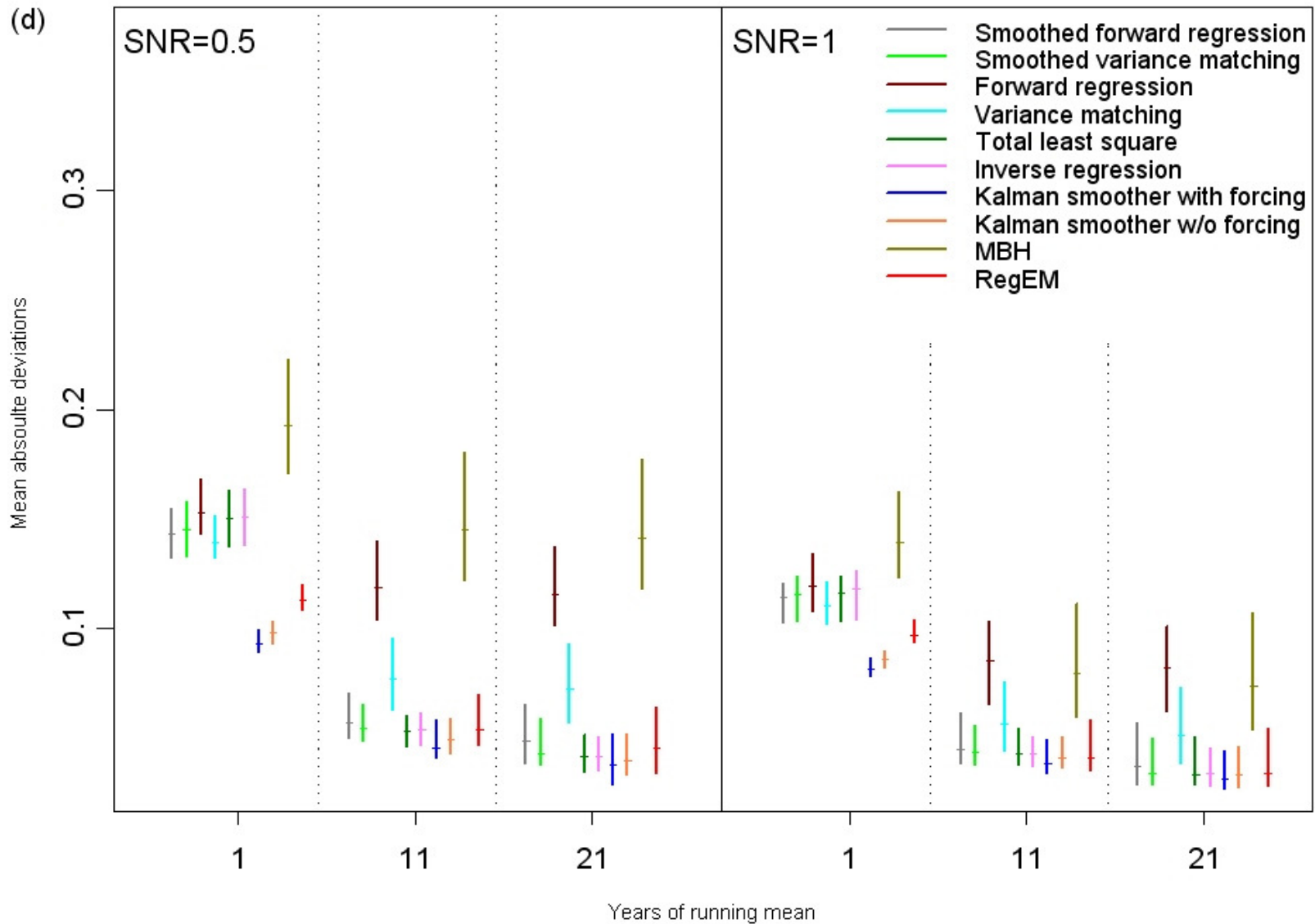


Some details

- We did this
 - Two different climate models
 - Two signal to noise ratios
 - White and red noise
 - Different sizes of proxy networks
 - Different lengths of calibration period
 - Effects of detrending prior to calibration
- We did NOT condition on a fixed proxy network
 - Repeated reconstruction 1000 times
 - A different proxy network at each time
 - A different realization of added noise each time

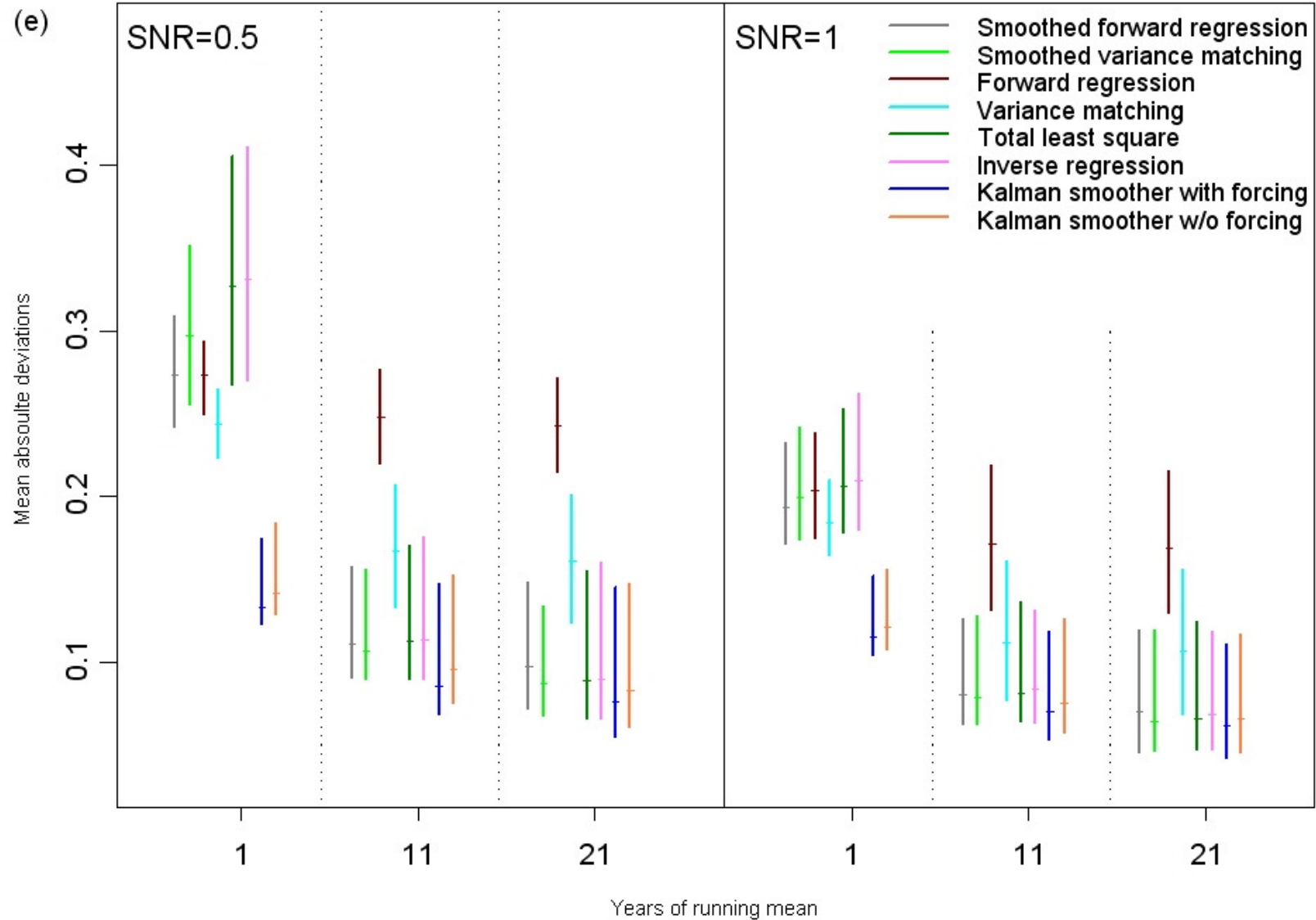
100 pseudo proxies – 1860-1970 calibration – mean abs deviation

ECHO-G



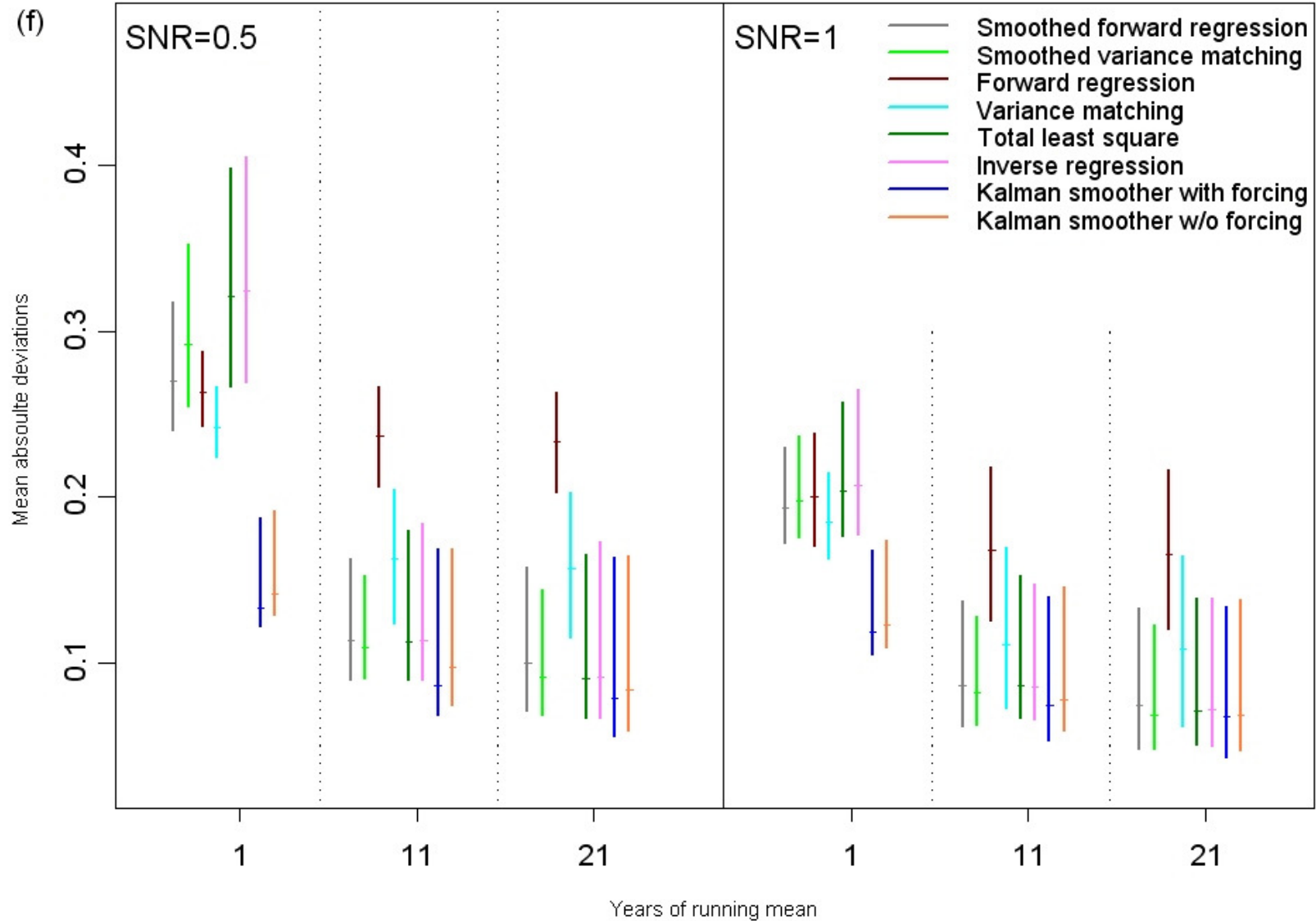
15 pseudo proxies – 1860-1970 calibration – mean abs deviation

ECHO-G



15 pseudo proxies – 1880-1960 calibration – mean abs deviation

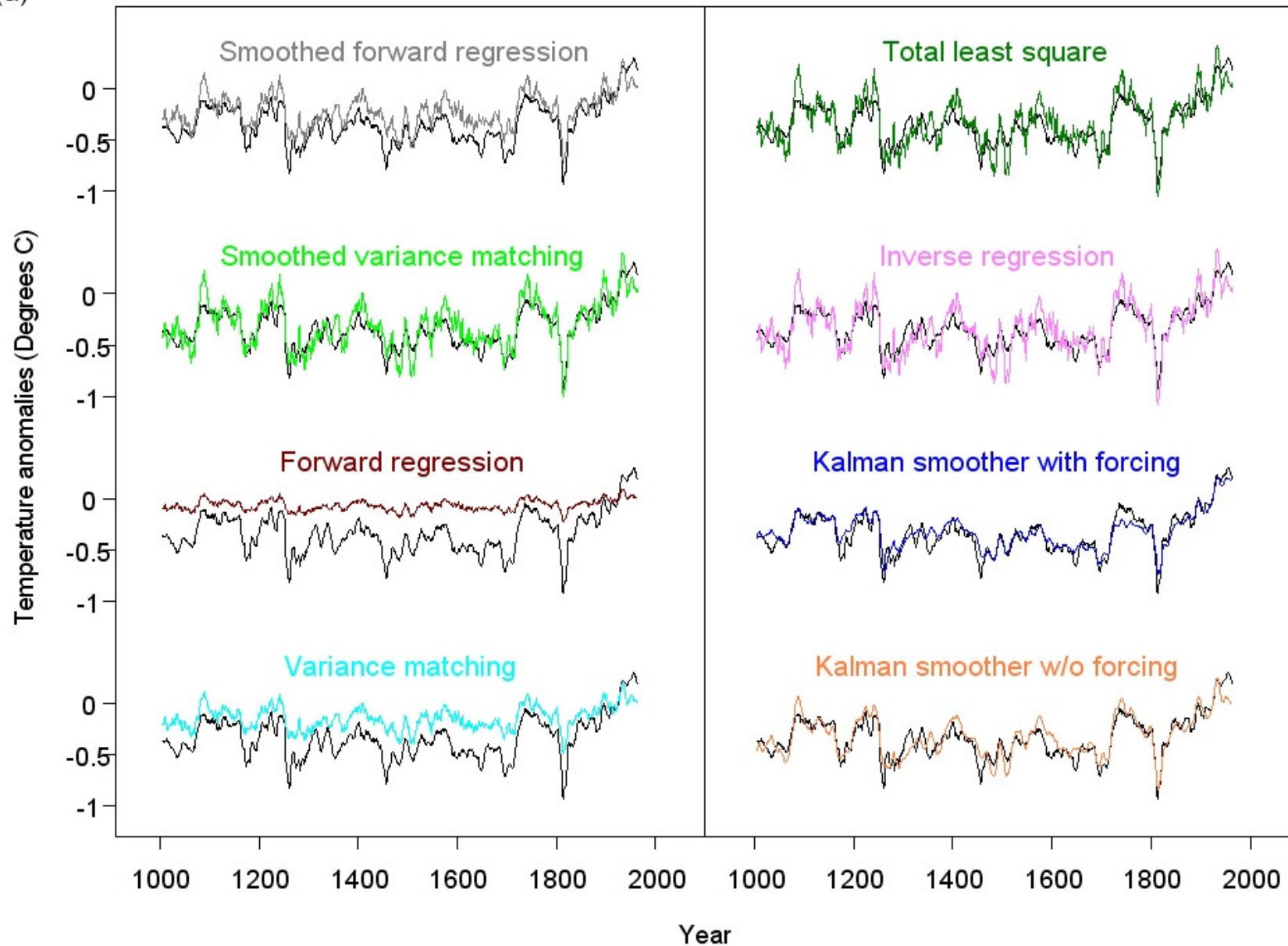
ECHO-G



15 point network – 11 year moving average – CSM - SNR = 0.5

One particular reconstruction from the sample of 100

(a)



Conclusions

- Some reconstruction methods to should be avoided
- Several other methods compare well, particularly at low frequency
- Assessment is not sensitive to which “reality” one compares against
- CFR methods do not have a particular predilection to producing hockey stick like reconstructions
- The new Mann et al CFR technique works well, but CFR does not appear to be consistently better than CPS
- Kalman filter approach appears to produce consistently good results across time scales
- There is probably greater sensitivity to the choice of proxy than to the choice of reconstruction method (see also Juckes et al, 2007, *Climate of the Past Discussions*, **vol 2**)

Photo: F. Zwiers



Two Extensions using Kalman smoother ...

State space model

$$P_t = \beta T_t + \varepsilon_t$$

$$T_t = \varphi T_{t-1} + \delta F_t + \omega_t$$

where

$$F_t = \alpha_{Sol} r(Sol_t) + \alpha_{Vol} r(Vol_t) + \alpha_G r(GHG_t) + \alpha_S r(Sul_t)$$

- Data assimilation (of sorts) using a climate model to assimilate proxy information and observed hemispheric temperatures
- Fitting the state space model provides estimates of coefficients and their uncertainty → allows “**detection**”
- Integrate climate model forward in time → allows “**prediction**”

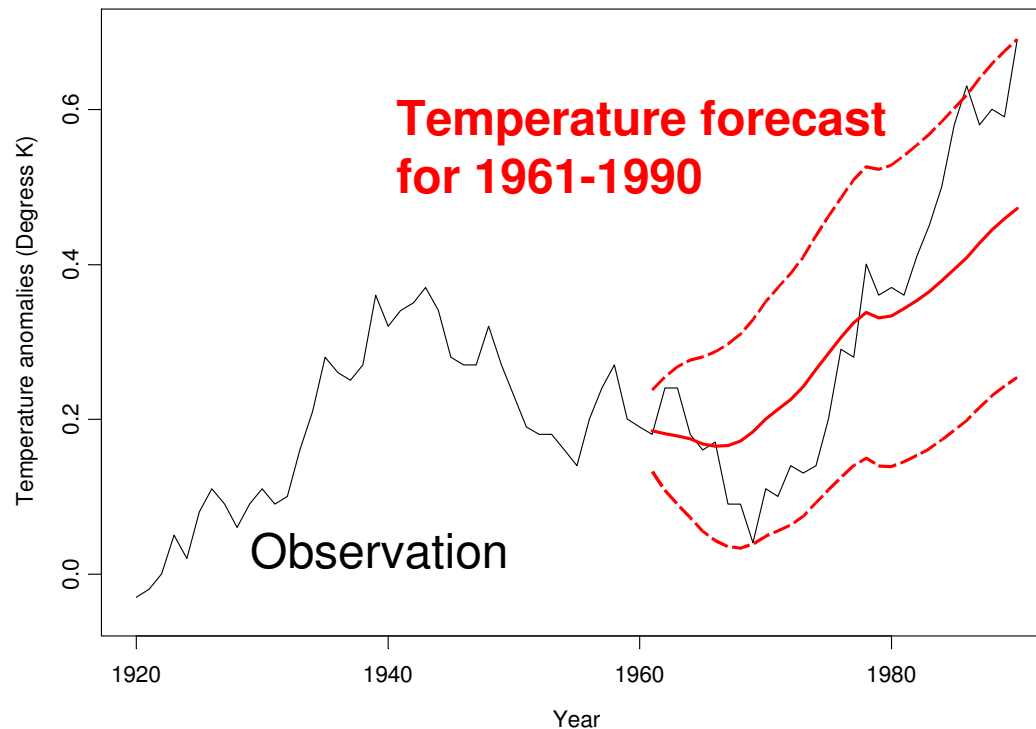
Results ...

Climate change detection:

- Using real-world proxy data (1510-1960):
 - Effect of greenhouse gas, aerosol and volcanic forcing has significantly affected our past climate.
 - Consistent with other detection studies.

Temperature prediction:

- Result with real-world proxy data (1510-1960):





The End

References

- Lee, T.C.K., F.W. Zwiers and M. Tsao, 2008: Evaluation of millennial proxy reconstruction methods. *Climate Dynamics*, **31**, 263-281, doi 10.1007/s00382-007-0351-9.
- Lee, T.C.K., M. Tsao, F.W. Zwiers, 2010: State-space model for proxy-based millennial reconstruction. *Canadian Journal of Statistics*, **38**, 488-505.