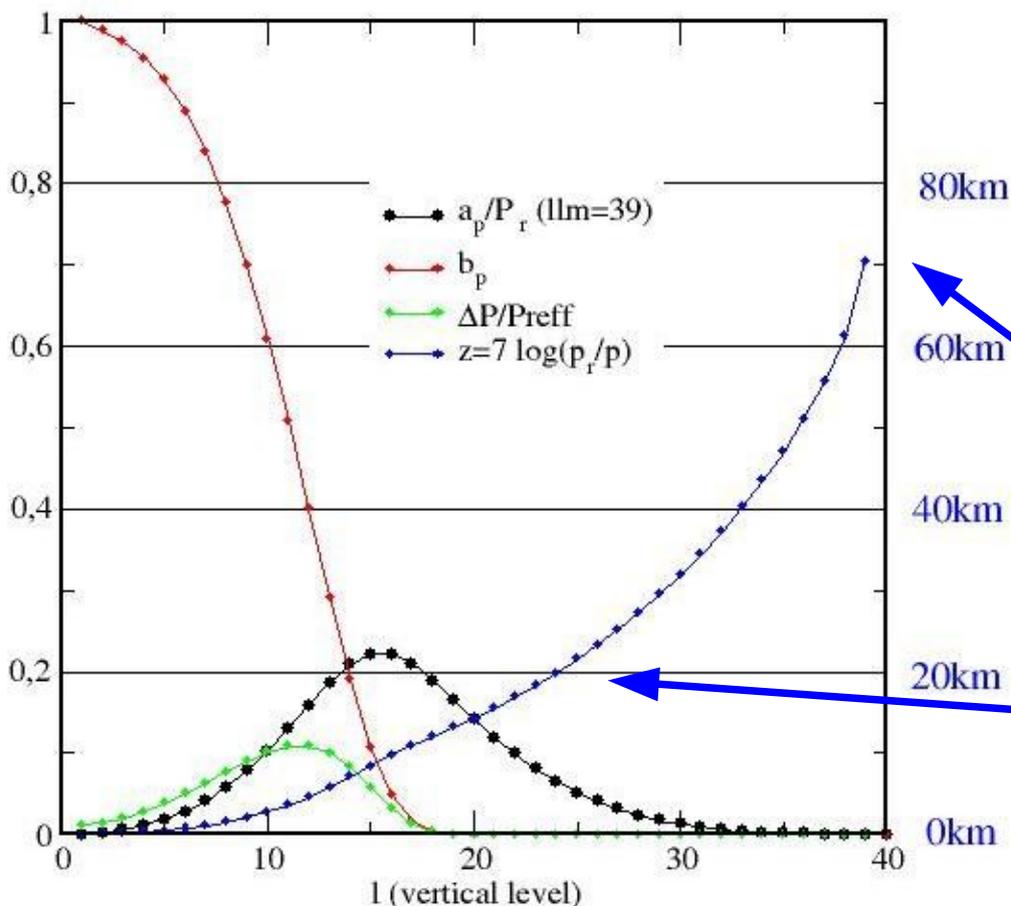


The tropical variability in the ongoing CMIP5 simulations at IPSL

Contacts for those diagnostics: F. Lott, P. Maury and L. Guez
LMD/IPSL, Ecole Normale Supérieure, Paris France

All the simulations done with the ESM IPSLCM5, include the stratosphere
The equilibrium pre-industrial 1000yrs, starting in 1800 control is done with the stratosphere. Historical runs and scenarios are also completed.



Ocean model: NEMO-LIM, $2^\circ \times 2^\circ \times 31$ levels with sea-ice model

Vegetation model: Orchidée

Coupling: OASIS from CERFACS

Atmospheric model: LMDz, 96x95x39

The experiments extend up to $Z_{top}=70\text{km}$, and includes orographic and non-orographic GWs parameterizations

The resolution in the low stratosphere is around 1.5-2km

Aerosols and volcanoes prescribed using INCA, Stratospheric Ozone prescribed via REPROBUS

Stratosphere and chemistry in LMDz: Lott et al. 2005, Jourdain et al. 2008
Performances in the midlatitudes: Nikulin and Lott (2010)

The tropical tropospheric variability in the ESM IPSLCM5

The model needs to have realistic tropospheric climate and variability (ENSO, MJO, and stratospheric PWs depend on these)
Also needed if one wishes to address which amount of waves needed for the QBO forcing are explicitly solved by the model

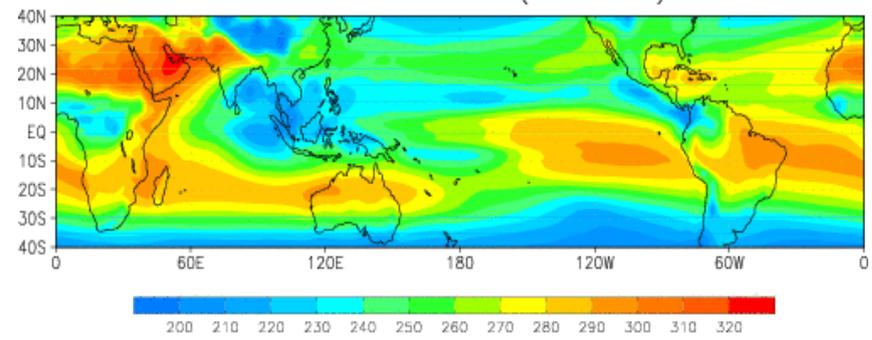
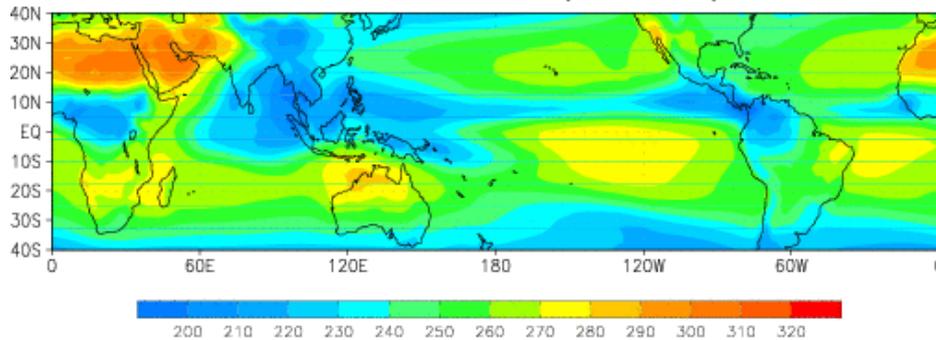
OLR diagnostics from the control run (1800-2350)

OLR NOAA (1979–2008)

OLR piControl2 (2200–2220)

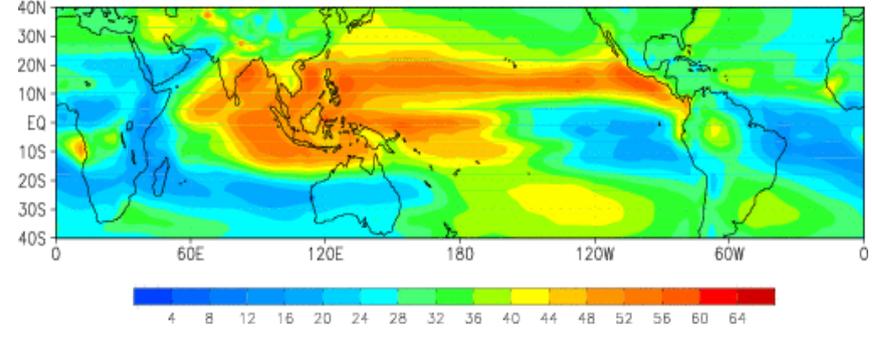
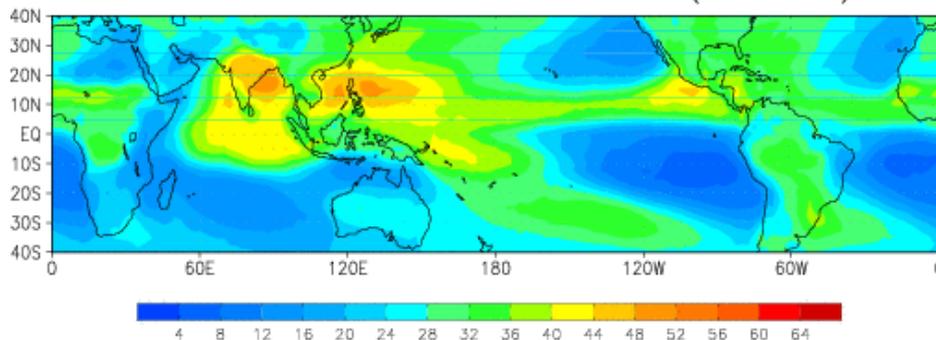
Mean Summer (MJJASO)

Mean Summer (MJJASO)



Standard Deviation Summer (MJJASO)

Standard Deviation Summer (MJJASO)



The tropical tropospheric variability in the ESM IPSLCM5

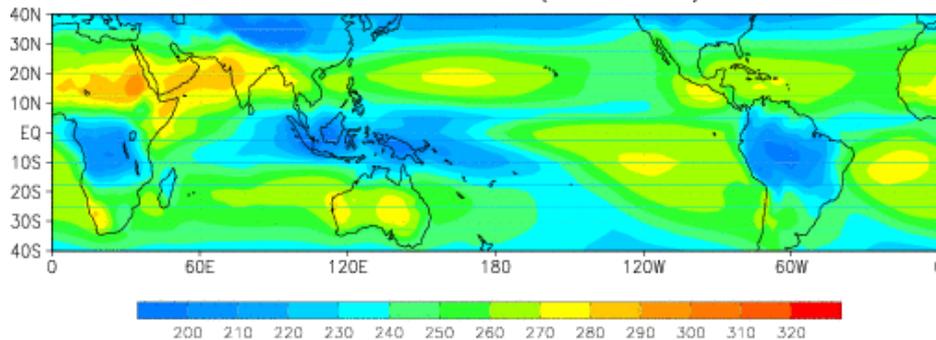
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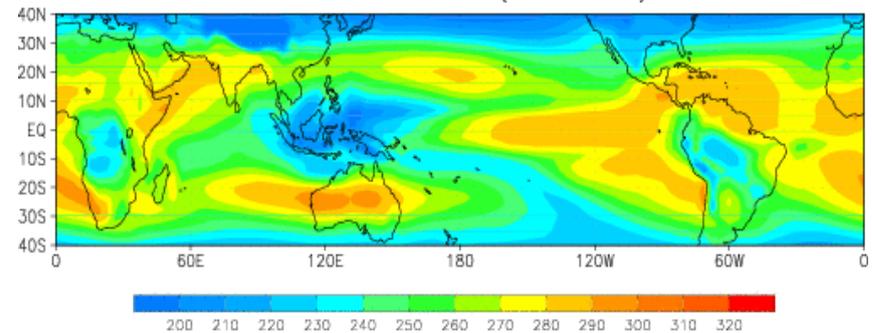
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OLR piControl2 (2200–2220)

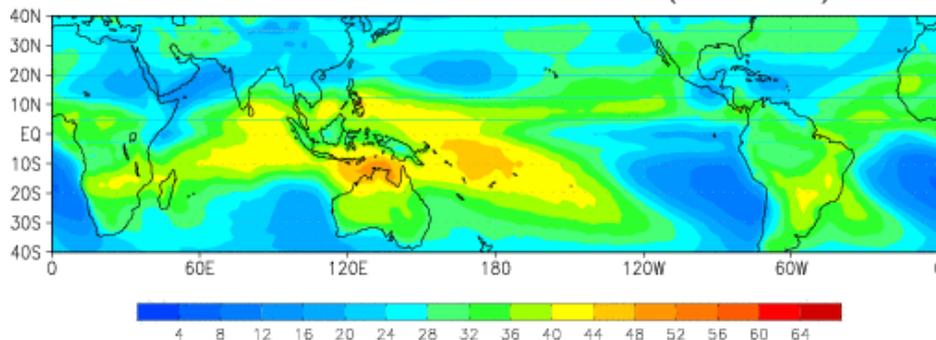
Mean Winter (SONDJF)



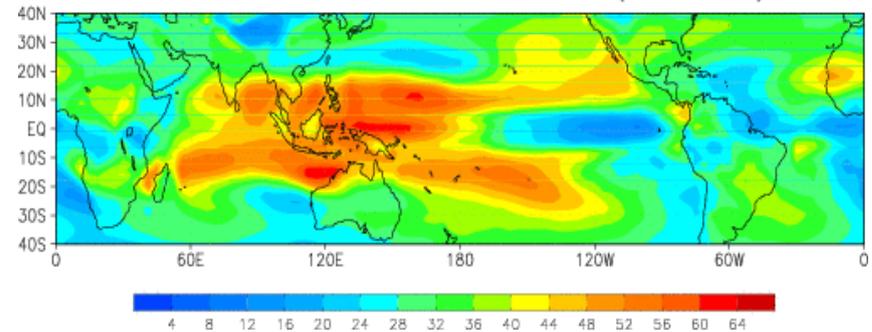
Mean Winter (SONDJF)



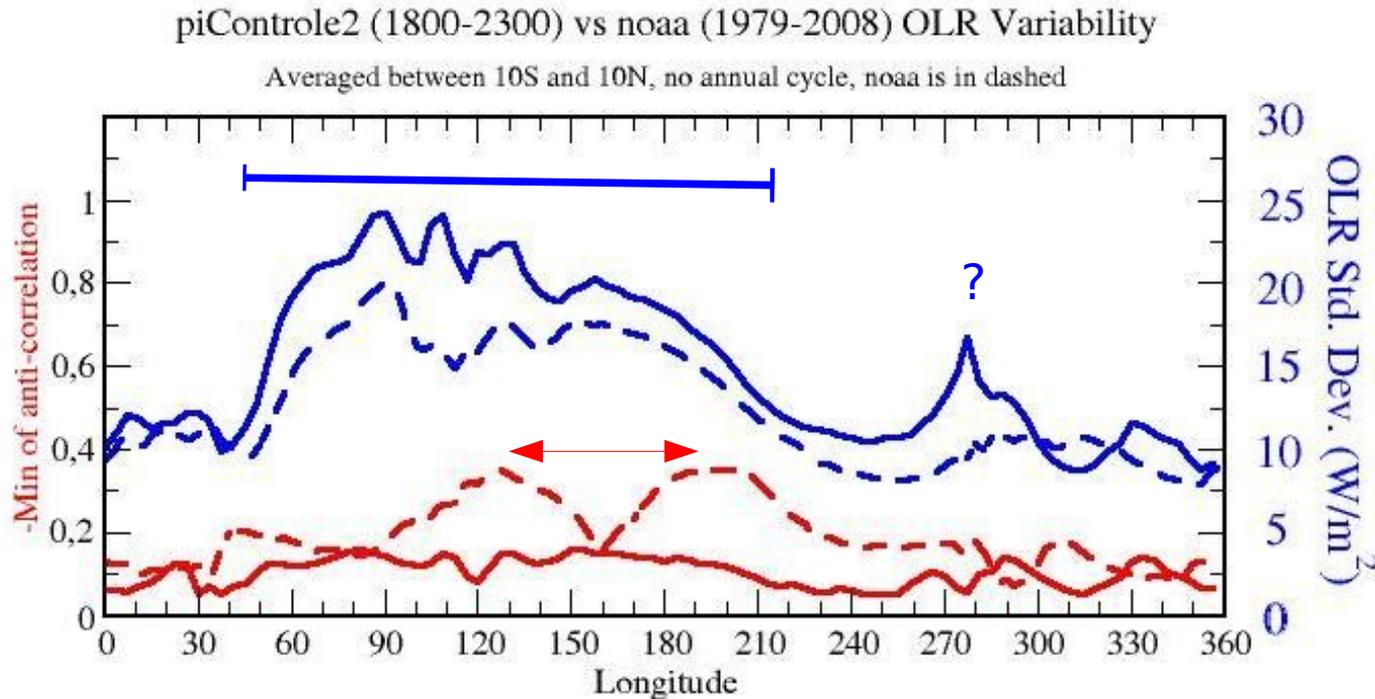
Standard Deviation Winter (SONDJF)



Standard Deviation Winter (SONDJF)



The tropical tropospheric variability in the ESM IPSLCM5

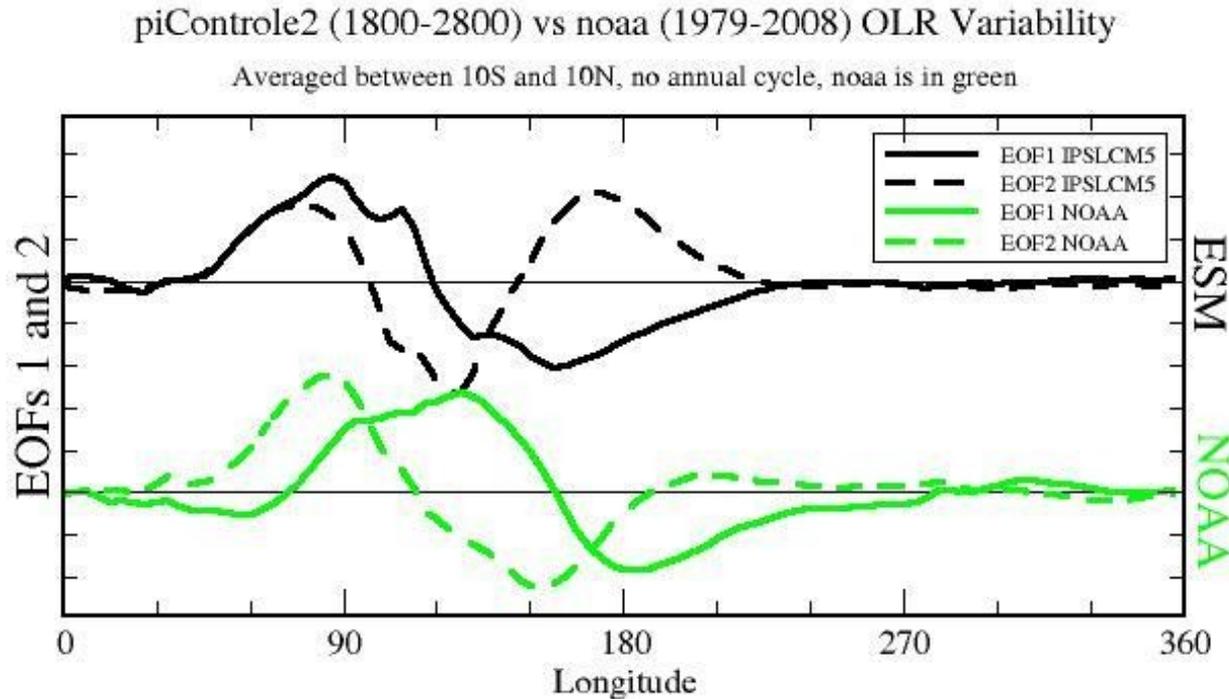


The model has enhanced variability from the Southern Indian ocean to the Mid-pacific, as in the Obs.

But a spurious peak of enhanced variability over Central and South America (?)

There is a significant underestimation of the anticorrelation between the maritime continent and the central pacific, remember that this anticorrelation is a signature of the Madden-Julian Oscillation (see the teleconnection arrow)

The tropical tropospheric variability in the ESM IPSLCM5

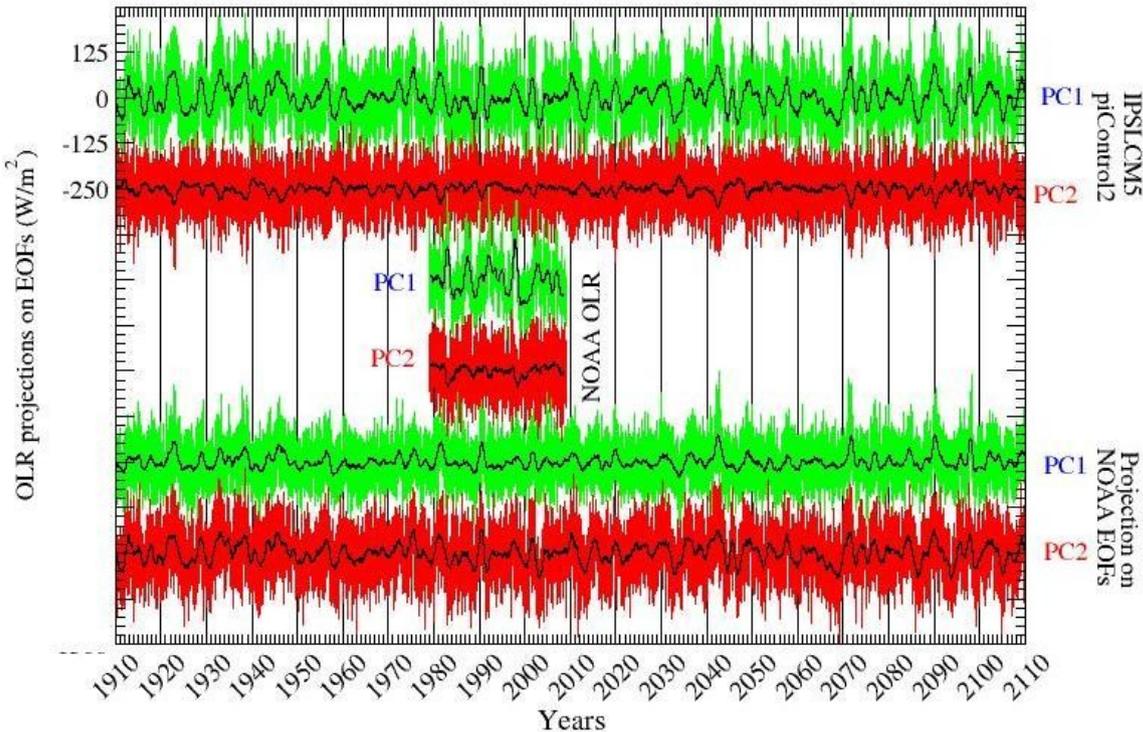


EOF1 in IPSLCM5 is more like the **EOF2** from observations
Both correspond to an excess in precip. over the western and central pacific;
and a deficit over the Equatorial Indian Ocean

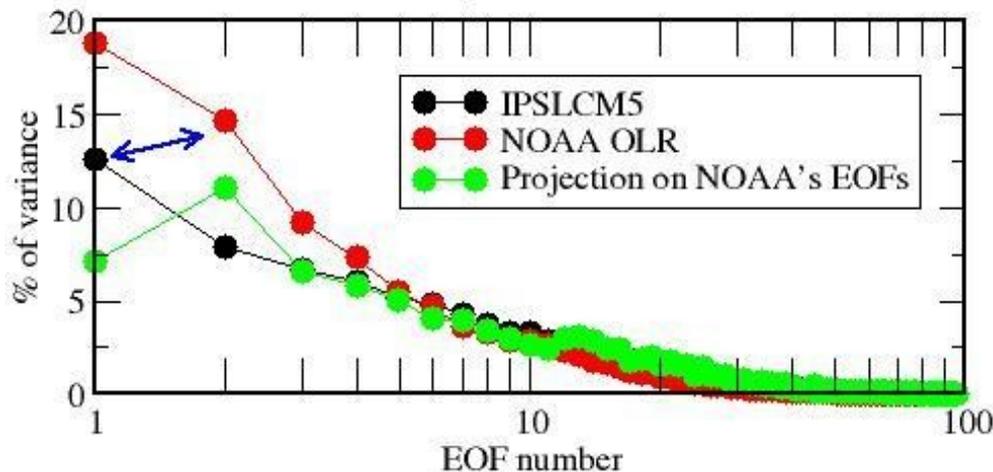
EOF2 in IPSLCM5 slightly reminiscent of like **EOF1** from observations
(but this is only true for that they are both associated with excess precipitation over
the maritime continent; for the western Indian ocean this is not clear at all!

The relative short scale of EOF2 (3 pronounced extrema) in IPSLCM5 call for
a more regional analysis.

The tropical tropospheric oscillations in the ESM IPSLCM5



EOF Spectrum of OLR



ENSO type:

PC1s in IPSLCM5 and NOAA shows more inter-Annual variability than PC2s

This is despite the fact that EOF1 in IPSLCM5 is More like EOF2 in NOAA!

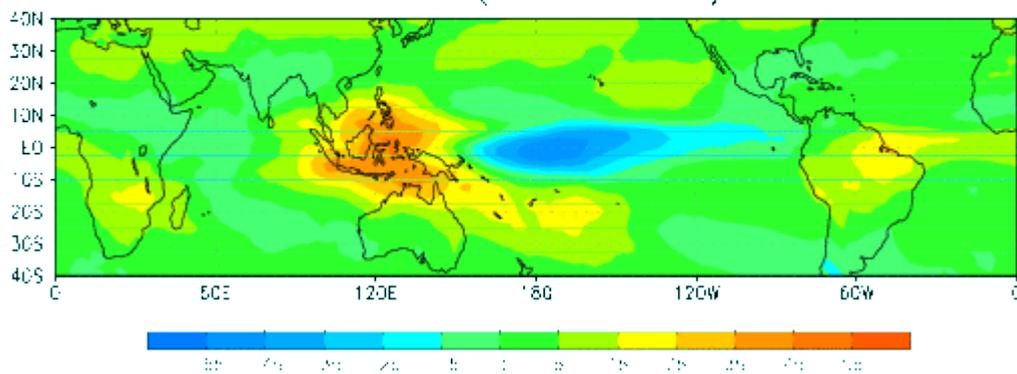
The Inter-Annual variability seems more confined to The western Pacific

The selected years are rather Insensitive if we choose EOF 1 from model or from observations to Attribute Nino years

The tropical tropospheric oscillations in the ESM IPSLCM5

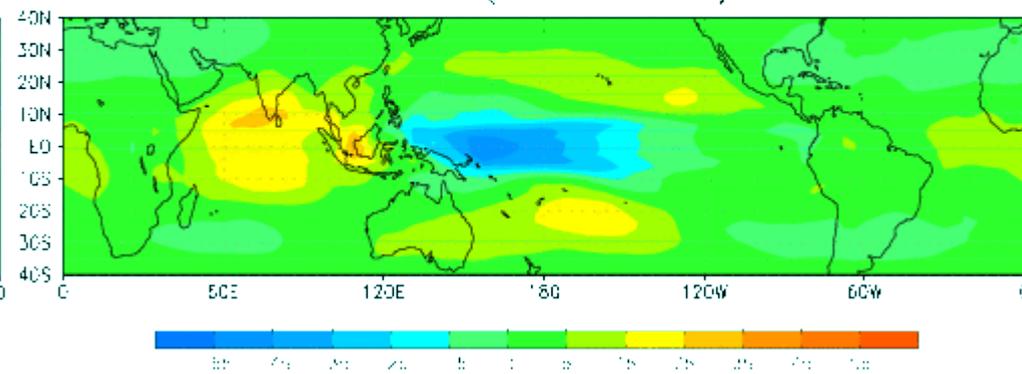
OLR NOAA (1979–2008)

Mean (NINO–NINA)

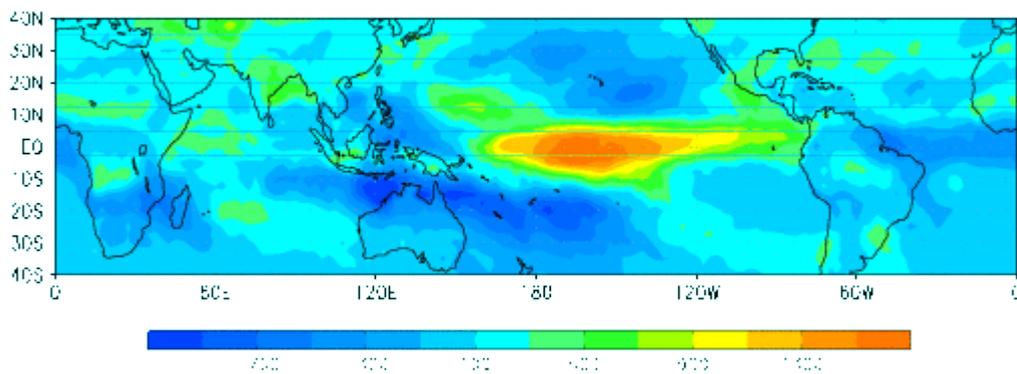


OLR piControl2 (1800–2000)

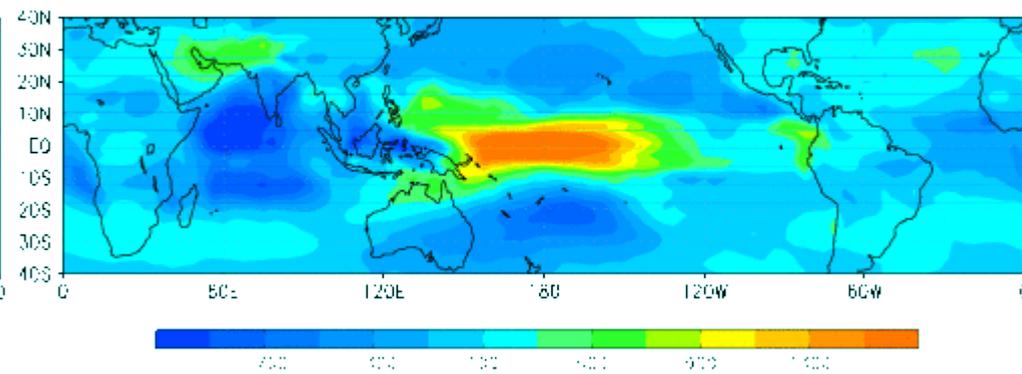
Mean (NINO–NINA)



shifts in Variance

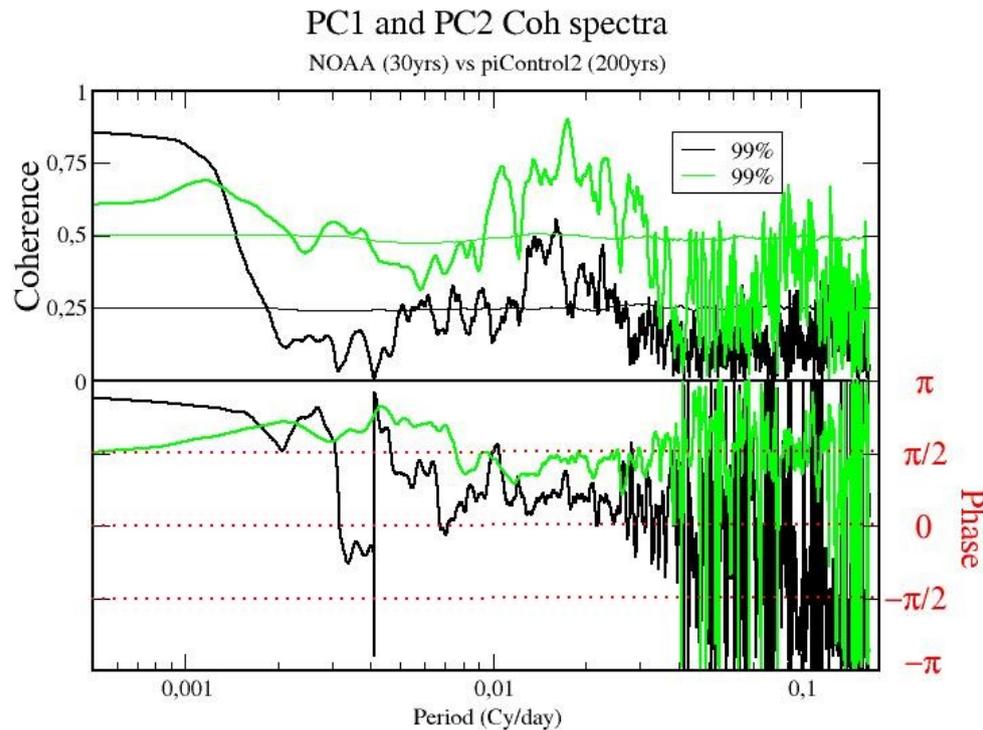


shifts in Variance



The tropical tropospheric oscillations in the ESM IPSLCM5

Coherency spectrum between
PC1 and PC2,
NOAA OLR Dashed (20 yrs)
PiControl2 (200yrs only, sorry!) Solid



MJO-type:

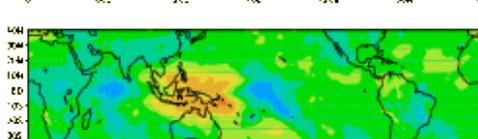
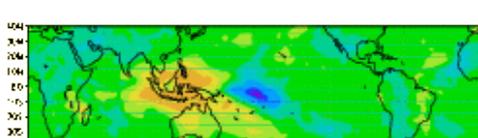
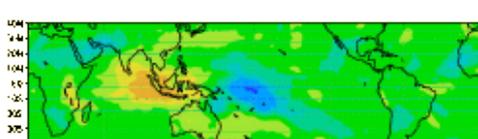
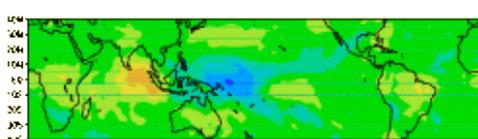
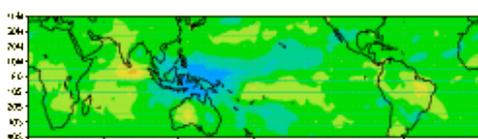
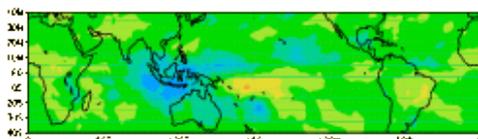
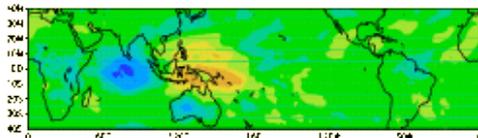
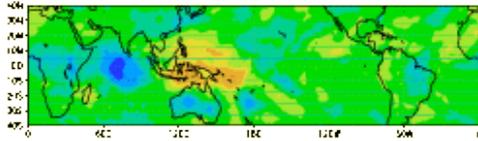
The intraseasonal variability is characterised in the NOAA OLR by the fact that the PC1 and PC2 signals are significantly coherent and in quadrature. This is almost absent from IPSLCM5

More precisely and in the IPSLCM5, the coherency is weak and the PC1 and 2 signals are almost in phase: the signal is more a standing oscillation than an Eastward propagating one.

The tropical tropospheric oscillations in the ESM IPSLCM5

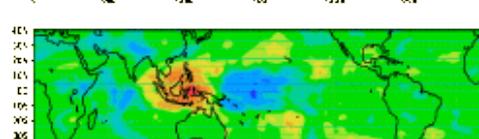
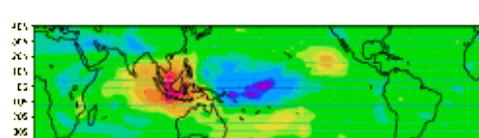
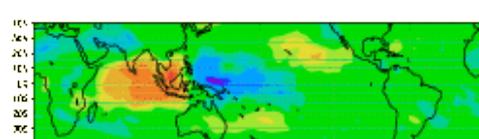
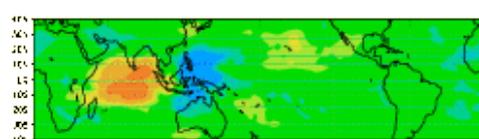
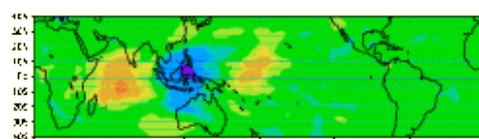
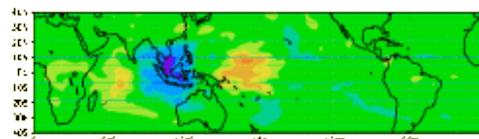
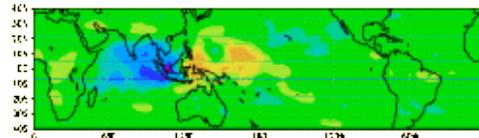
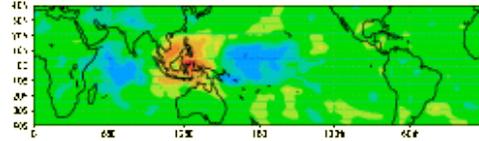
NOAA-OLR

NOAA OLR (1979-2008) 14 Cases



π ctl2-OLR

MJO in π ctl2 OLR (1800-2799) 68 Cases



Composite maps of OLR keyed to the amplitude of the (PC1,PC2) vector, filtered in the IS band.

Only events lasting more than 30 days are kept

Composite MJO out of π ctl2 (68 cases out of 1000yrs!).

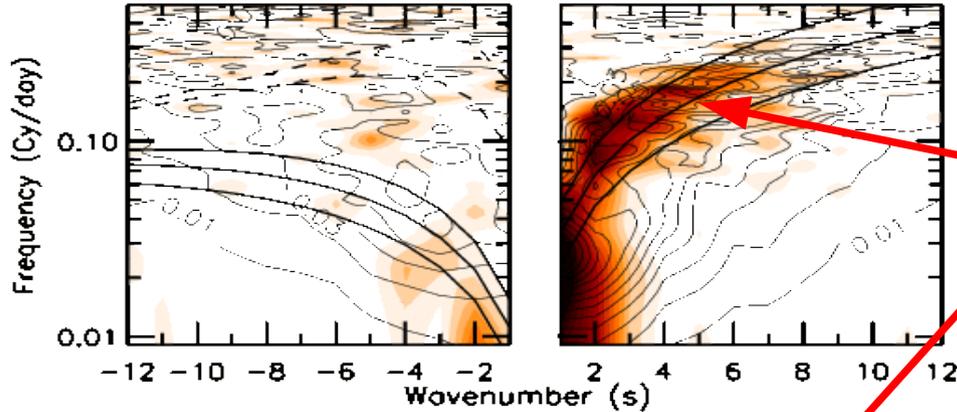
Long lasting ones propagate properly, but there are very few!



The tropical waves in the ESM IPSLCM5

U Sym vs Precips Sym [-10S,10N], NCEP2 and GPCP 1997-2008
Precips and Coh**2 of {Precip vs U 250hPa}

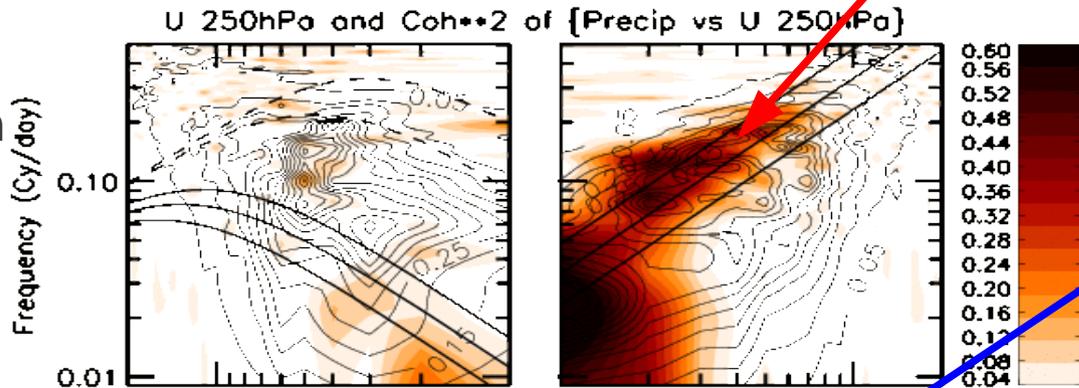
Precips
from
GPCP



The tropospheric convectively
coupled equatorial waves
(here of Kelvin wave type)

are not so distinct from

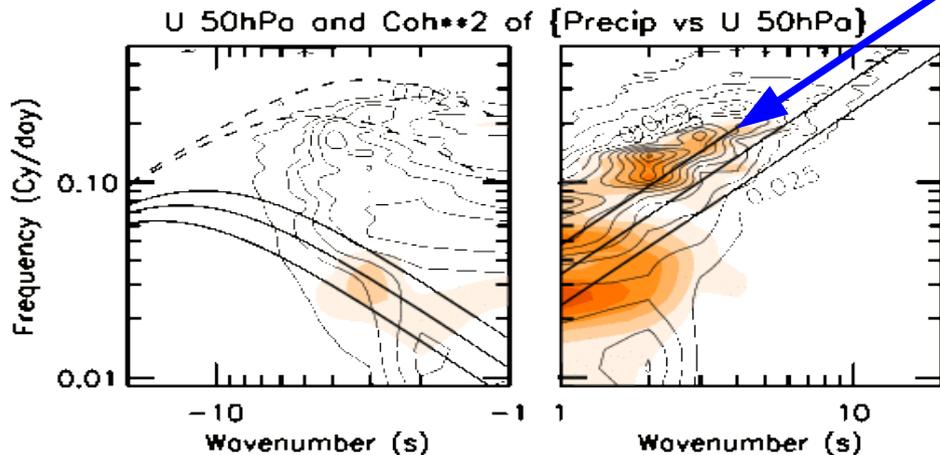
U250hPa
from
NCEP



The freely propagating
Stratospheric waves

Is it critical for the model
Stratospheres since
CCEWs are known to
be weak in models?

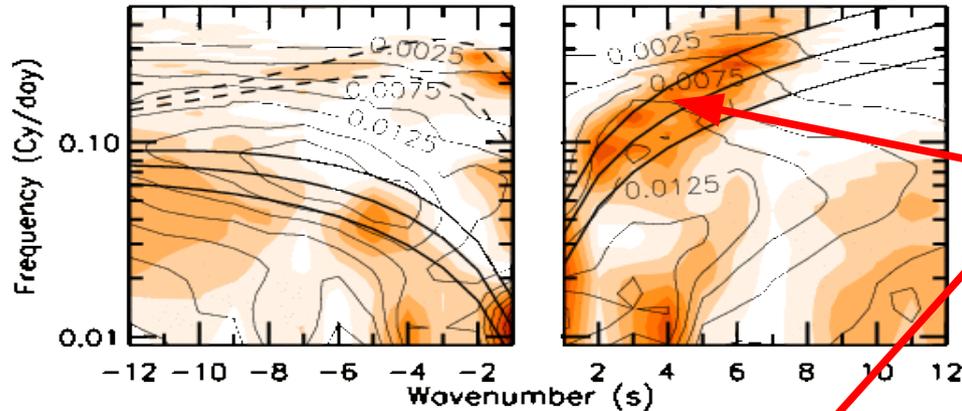
U50hPa
from
NCEP



Redish shadings are for coherencies
between speeds and precipitations

The tropical waves in the ESM IPSLCM5

U Sym vs Precips Sym [-10S,10N], piControl2 1800-1999
Precips and Coh**2 of {Precip vs U 250hPa}

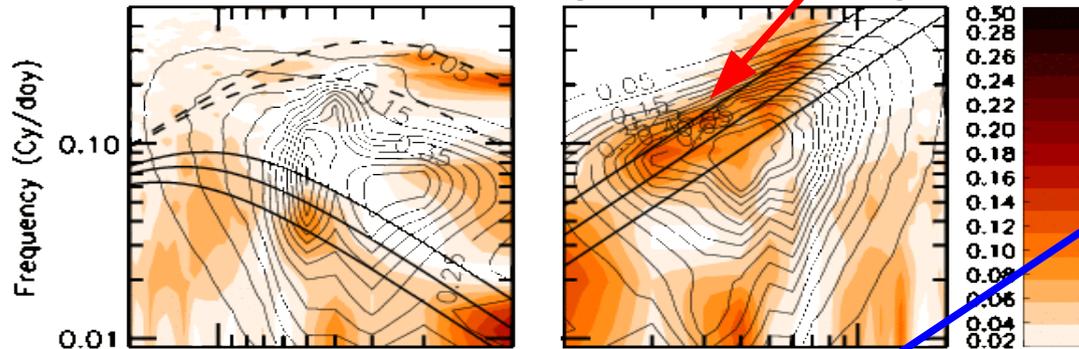


Precips
from
 π Ctl2

The tropospheric convectively coupled equatorial waves (here of Kelvin wave type)

are not so distinct from

U 250hPa and Coh**2 of {Precip vs U 250hPa}

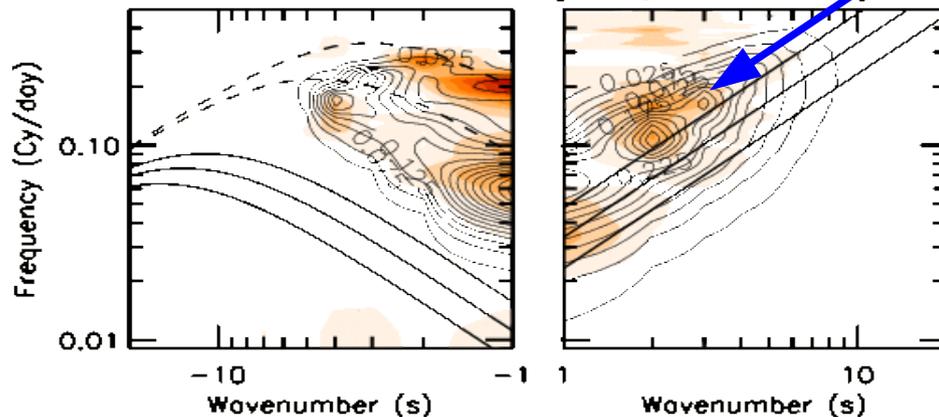


U250hPa
from
 π Ctl2

The freely propagating Stratospheric waves

Is it critical for the model Stratospheres since CCEWs are known to be weak in models?

U 50hPa and Coh**2 of {Precip vs U 50hPa}



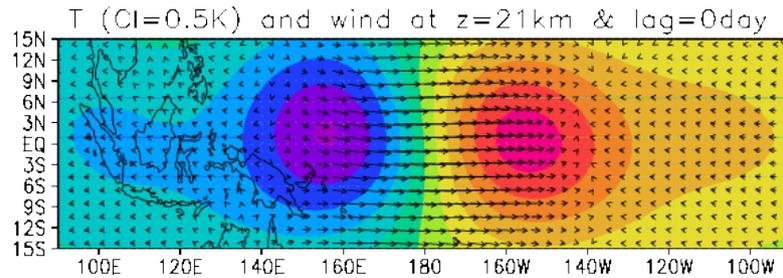
U50hPa
from
 π Ctl2

Redish shadings are for coherencies between speeds and precipitations

The tropical waves in the ESM IPSLCM5

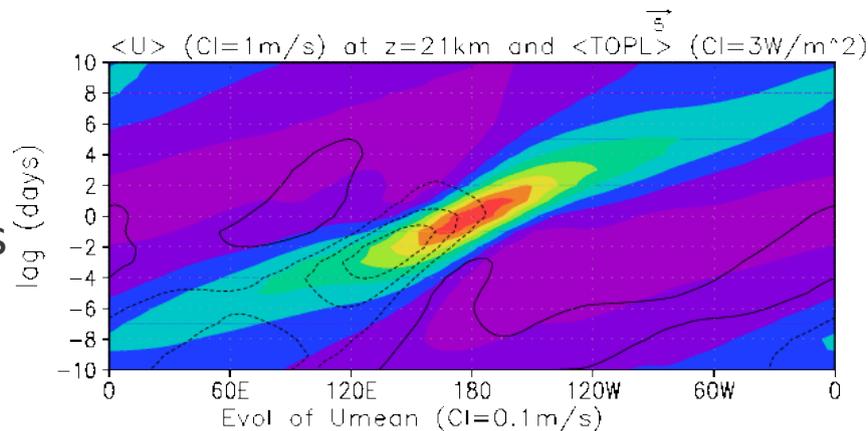
Wave-Mean flow interactions

Composite kelvin in piCtI2 (1800–2000)



T is in color,

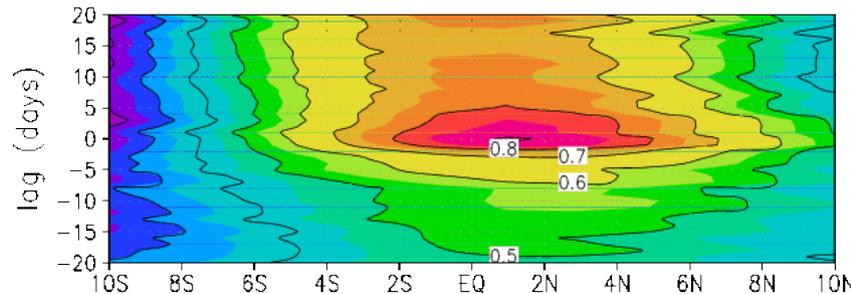
Composite analysis illustrates better the structure of the waves



U is in color,
OLR in contours

(here at 50hPa, except for the OLR)

Zonal mean
of U



Evolution of the zonal mean wind during the passage of the Kelvin waves (wave-mean flow interaction clearly visible here)

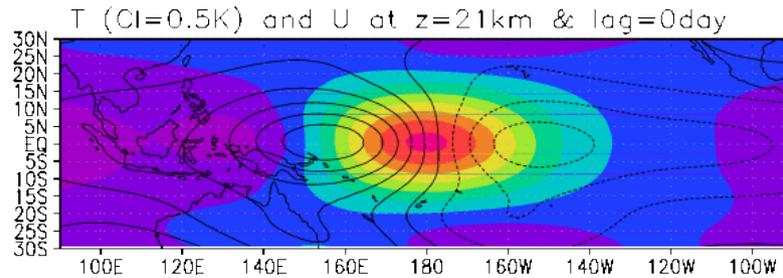
Amplitude comparable to those documented in Lott et al. (2009)

The tropical waves in the ESM IPSLCM5

Dehydration (here more a moistening!)

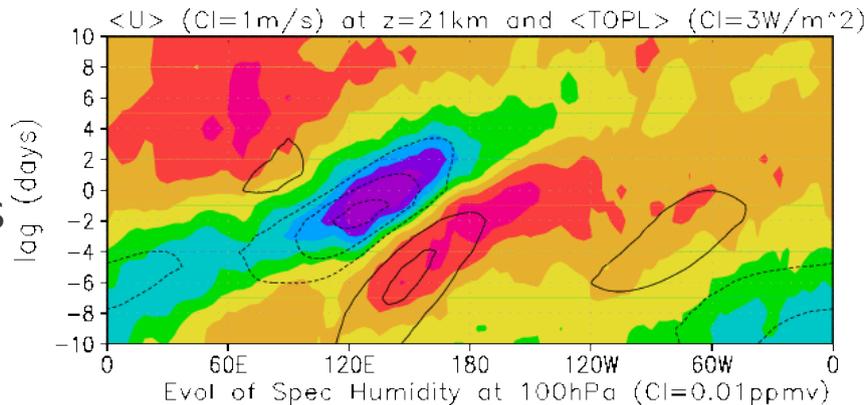
Composite keldhy in piCtI2 (1800–2000)

T is in color,
U in shades



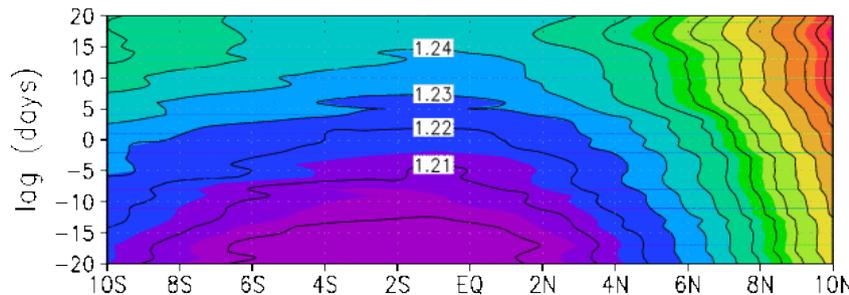
Composite analysis
illustrates better the
structure of the waves

U is in color,
OLR in contours



(here at 50hPa, except
for the water vapour which
is at 100hPa)

Zonal mean
of U



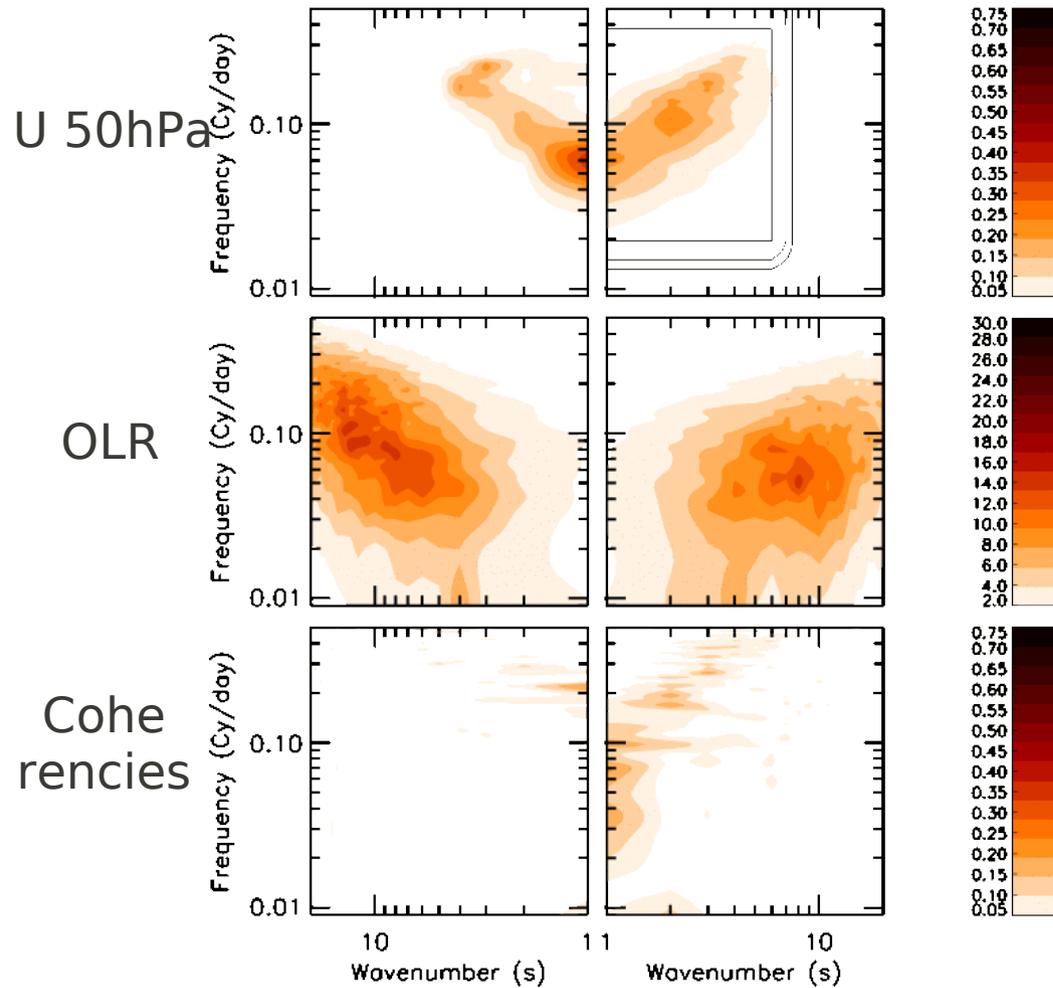
Evolution of the water vapor
mixing ratio at 100hPa
(stratospheric moistening
clearly visible here)

Amplitude comparable to those documented in Lott et al. (2009)

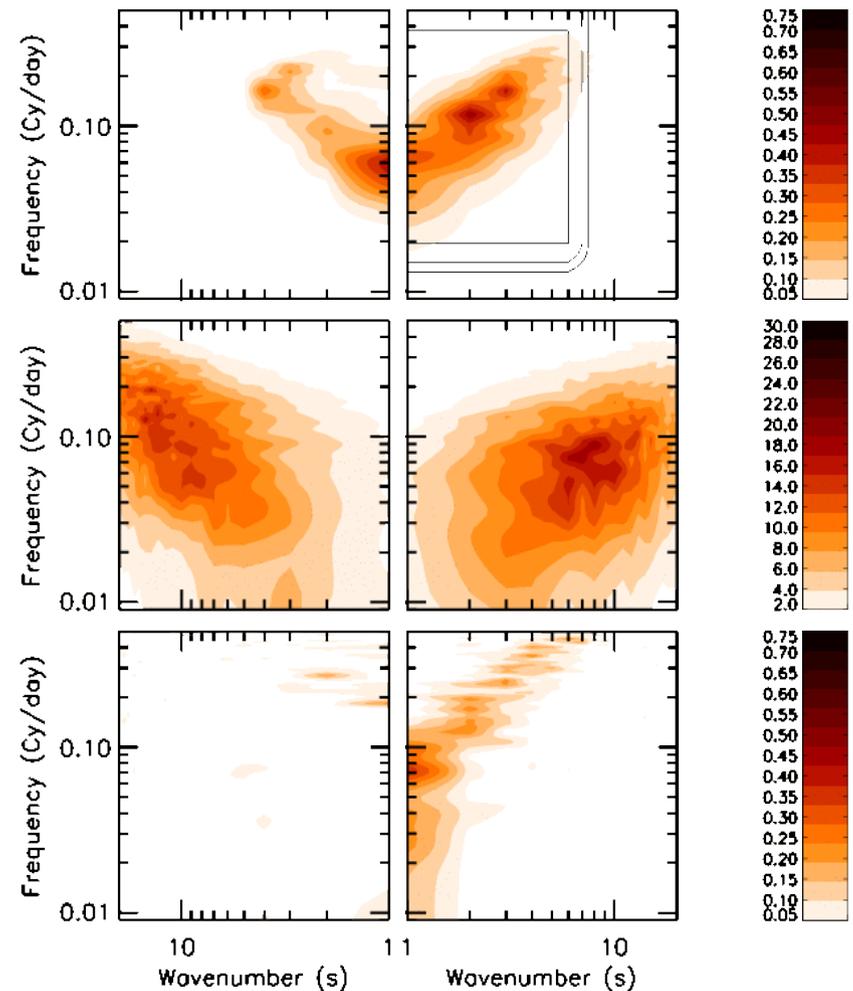
The tropical waves in the ESM IPSLCM5

Weak sensitivity to ENSO! (is there is an ENSO to QBO relation?)

U Sym and OLR Sym, PiCtI2 Westward 20 ENSO- of 1800-2799



U Sym and OLR Sym, PiCtI2 Westward 20 ENSO+ of 1800-2799



The tropical variability in the ongoing CMIP5 simulations at IPSL

Contacts for those diagnostics: F. Lott, P. Maury and L. Guez
LMD/IPSL, Ecole Normale Supérieure, Paris France

Take care when you look at the impact of El-Nino in Southern America or elsewhere; the principal center of variability is very significantly shifted westward; this can affect remote impacts via Rossby waves propagation from this center.

Can we do regional modeling over south America from this runs?