Understanding Climate Related Process in the Southwestern Atlantic Ocean

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- Mesoscale eddies in the SWA Ocean (CIMA, LOCEAN)
- Monitoring the cold route in the SWA Ocean (LOCEAN, SHN)
- Monitoring the ACC transport through Drake Passage (LOCEAN, SHN, CIMA)

Mesoscale eddies in the Southwestern Atlantic Ocean

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Objectives:

To study the relationship between eddies distribution and main flow in the Southwestern Atlantic



The Southwestern Atlantic region



The cold and low-salinity waters transported poleward by the Malvinas Current encounter the warm and salty subtropical waters transported by the Brazil Current at about 38°S.



Mean fields of several variables reveal clear patterns

Mean fields are very different from instantaneous fields:



Satellite radar altimetry data do not provide instantaneous maps of SSH but do reveal mesoscale eddies – and without clouds



SSH map representative of a seven day period obtained from the interpolation in space and time of 4 different missions (Jason 1, ERS, GFO and T/P). Black vectors represent geostrophic velocities.

Eddies play a crucial role in the region:



[[]Barré et al., 2006]

- Eddies can transport momentum, heat, potential vorticity but also nutrients, chlorophyll and zooplankton
- Eddies thus have a direct implication for the role of the oceans in the global heat balance.

Much of the mesoscale variability outside of the tropics consists of nonlinear eddies.

 Far from the equatorial band (+-20°), the largest portion of the mesoscale variability observed is due to nonlinear interaction of eddies, not to Rossby waves.

[Chelton et al., 2007] - nondia



- nondispersive baroclinic Rossby waves
- large-scale, long-lived eddies
- short-scale, long-lived eddies

Eddies interact with the Malvinas Current

Malvinas Current transport at 41°S from altimetry (black line) and in-situ data (red and blue):



Methodology used: Okubo-Weiss algorithm

(i) Geostrophic velocities & derivatives:

(ii) The following parameters are estimated

Shearing deformation rate: $s_s = g_{21} + g_{12}$,

Stretching deformation rate $S_n = g_{11} - g_{22}$,

Total deformation rate :
$$s = \sqrt{s_s^2 + s_n^2}$$

Divergence: $\psi = g_{11} + g_{22}$. Vorticity: $\omega = g_{21} - g_{12}$,

Spatial distribution of eddies



- Large number of eddies in the regions with high mesoscale activity: in the Brazil-Malvinas Confluence region and along the Subantarctic Front.

- Regions corresponding to the Zapiola Rise, the Malvinas Current and north of the Subtropical Front present a minimum number of eddies.

Trajectories of long-lived (> 6 weeks) eddies



Trajectories of the long lived C & A eddies are similar. Cyclonic eddies are found more often in the Zapiola Rise region compared with anticyclonic ones.

Radius (km)



The spatial distribution of the mean radius of the eddies shows an homogeneous distribution pattern. The number, mean radius and spatial distribution of these variables does not varies significantly for the total cyclonic or anticyclonic eddies. Similar results are found in other regions of the ocean (Chaigneau at al, 2009)

Interannual variability in the SWA (whole region)



Anticyclonic



Interannual variability in the Overshoot region



Interannual variability in the Overshoot region





Are eddies in the Overshoot region the main forcing of the circulation around ZR?

Conclusions

- Eddies detected with the Okubo-Weiss method in the SWA originate mainly in the Brazil-Malvinas Confluence region and along the Subantarctic Front, where the EKE is higher.
- Eddies circulate around and do not enter the Zapiola Rise (ZR) region. This observation is in agreement with the eddy-driven circulation theory [e.g. Dewar, 1998].
- It is suggested that the interannual variability of the transport associated to the ZR may be associated to the number of eddies in the Confluence region.

Monitoring the cold route in the SWA Ocean

- 1- monitoring transport of the Malvinas Current (Moorings + altimetry)
- 2- documenting water mass transformation processes in the Argentine Basin (Argo floats)
- 3- monitoring transport and water masses variability through Drake Passage (Hydrography, Moorings, altimetry)



Malvinas Current : 15 year transport time series



Mean 34.3 Sv, std 7.4 Sv

> good correlation (> 0.7 [95% CL=0.48] for each period) between transport computed from currentmeter data (alone) and transport time series from altimetry

Spadone and Provost, 2009. JGR-Oceans

Time evolution of the spectral content



> modulations at all frequencies

> apparition of a signal around the annual period after 2000 (maximum in 2001-2006 seems to be decreasing afterwards)

Variations : local or remote causes ?



Surface geostrophic velocity anomalies between 2 f/H contours corresponding to the core of the MC





DRAKE Moorings



Jason track # 104 Eastern side of Shackleton Fracture Zone West Scotia Ridge Yaghan Basin - Ona Basin

- Feb 2006: M1 to M10 deployed belowJason track # 104 at crossover points.M2, better resolution in the North
- 8 moorings recovered in April 2008
 M2 and M8 lost
- 5 moorings M1-M5 redeployed in April 2008
- All 5 moorings recovered March-April 2009

Over 3 years in Yaghan Basin

2 years in Ona Basin

Data gathered during DRAKE Project

4 Hydrographic cruises from RV POLARSTERN

3 with tracers (V. Garçon, O. Huhn, H. de Baar) and LADCP

2005 April E. Fahrbach

21 CTD stations no LADCP no tracers Precise bathymetry

2006 January-February 2 sections along track 104

51 CTD/LADCP/tracers stations southward 43 CTD/LADCP/tracers stations northward





Data gathered during Drake Project

4 Hydrographic cruises from RV POLARSTERN

2008 - ZERO-DRAKE (IPY) E. Fahrbach - Feb - April 31 CTD/ 29 LADCP stations (#104)



2009 March-April (poor weather) 25 CTD/LADCP stations southward (#104) 17 CTD/LADCP stations northward on the SFZ (Jason track #28)



Two crossings of the Drake Passage in austral summer 2006



C. Provost, <u>A. Renault</u>, N. Barré, N. Sennéchael, V. Garçon, J. Sudre and O. Huhn

Spatial context: Jan-Feb 2006



0.075

0.1

02

0.15

0.3

0.5

0.7

1

1.5

Southward Transect: 51 stations

320 300





- Fronts- branches eddies •
- Signature in velocities, • hydrography, altimetry

Northward Transect: 43 stations

300





- Time evolution between the 2 transects
- Fronts- branches eddies
- Signature in velocities, hydrography, altimetry



LADCP Data

- ✓ The error bar is small over the entire section (less than 3 cm/s) with few exceptions
- ✓ Velocity structure in narrow bands
- ✓ The fronts are well defined in the across track velocities
- ✓ Several eddies are observed
- ✓ Evolution within three weeks

Position and strength of the PF and SACCF are modifiedSmaller westward velocities



Transport estimates



Differences in properties

- High differences between the 10-day apart sections
 - Below 3000 m:
 - ΔT=0.2°C,
 - ΔS=0.01,
 - Δγⁿ=0.03 kg/m³,
 - ΔOxy=10 µmol/kg
- Only part attributable to displacements of fronts and eddies
- Spatial heterogeneity of water properties upstream
 - PF: heterogeneity upstream, funnelling of the flow due to topographic constraints
- Strong differences in water properties on short time scales

Caution when interpreting different sections in terms of climatic signals



Conclusions

- Fine horizontal resolution of the hydrographic data and high accuracy of the data
- 2 sections repeated within 3 weeks
 - Unique opportunity to document full depth variability at 10day interval
- Contributions of fronts and eddies to total transport changed
 - No strong modification of the transport (about 10%)
- Strong differences in water properties



- Frontal or eddy displacements
- Spatial differences of properties upstream the section